



Influence of aerosols and surface reflectance on satellite NO₂ retrieval: seasonal and spatial characteristics and implications for NO_x emission constraints

J.-T. Lin¹, M.-Y. Liu¹, J.-Y. Xin², K. F. Boersma^{3,4}, R. Spurr⁵, R. Martin^{6,7}, and Q. Zhang⁸

¹Laboratory for Climate and Ocean-Atmosphere Studies, Department of Atmospheric and Oceanic Sciences, School of Physics, Peking University, Beijing 100871, China

²State Key Laboratory of Atmospheric Boundary Layer Physics and Atmospheric Chemistry, Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing, 100029, PR China

³Royal Netherlands Meteorological Institute, De Bilt, the Netherlands

⁴Meteorology and Air Quality department, Wageningen University, Wageningen, the Netherlands

⁵RT Solutions Inc., Cambridge, MA 02138, USA

⁶Department of Physics and Atmospheric Science, Dalhousie University, Halifax, Nova Scotia, Canada

⁷Atomic and Molecular Physics Division, Harvard – Smithsonian Center for Astrophysics, Cambridge, Massachusetts, USA

⁸Ministry of Education Key Laboratory for Earth System Modeling, Center for Earth System Science, Tsinghua University, Beijing, China

Correspondence to: J.-T. Lin (linjt@pku.edu.cn)

Received: 25 March 2015 – Published in Atmos. Chem. Phys. Discuss.: 29 April 2015

Revised: 4 August 2015 – Accepted: 22 September 2015 – Published: 9 October 2015

Abstract. Satellite retrievals of vertical column densities (VCDs) of tropospheric nitrogen dioxide (NO₂) normally do not explicitly account for aerosol optical effects and surface reflectance anisotropy that vary with space and time. Here, we conduct an improved retrieval of NO₂ VCDs over China, called the POMINO algorithm, based on measurements from the Ozone Monitoring Instrument (OMI), and we test the importance of a number of aerosol and surface reflectance treatments in this algorithm. POMINO uses a parallelized LIDORT-driven AMFv6 package to derive tropospheric air mass factors via pixel-specific radiative transfer calculations with no look-up tables, taking slant column densities from DOMINO v2. Prerequisite cloud optical properties are derived from a dedicated cloud retrieval process that is fully consistent with the main NO₂ retrieval. Aerosol optical properties are taken from GEOS-Chem simulations constrained by MODIS aerosol optical depth (AOD) data. MODIS bi-directional reflectance distribution function (BRDF) data are used for surface reflectance over land. For the present analysis, POMINO level-2 data for 2012 are aggregated into monthly means on a 0.25° long. × 0.25° lat. grid.

POMINO-retrieved annual mean NO₂ VCDs vary from 15–25 × 10¹⁵ cm⁻² over the polluted North China Plain (NCP) to below 10¹⁵ cm⁻² over much of western China. Using POMINO to infer Chinese emissions of nitrogen oxides leads to annual anthropogenic emissions of 9.05 TgN yr⁻¹, an increase from 2006 (Lin, 2012) by about 19%. Replacing the MODIS BRDF data with the OMLER v1 monthly climatological albedo data affects NO₂ VCDs by up to 40% for certain locations and seasons. The effect on constrained NO_x emissions is small. Excluding aerosol information from the retrieval process (this is the traditional “implicit” treatment) enhances annual mean NO₂ VCDs by 15–40% over much of eastern China. Seasonally, NO₂ VCDs are reduced by 10–20% over parts of the NCP in spring and over northern China in winter, despite the general enhancements in summer and fall. The effect on subsequently constrained annual emissions is between –5 and +30% with large seasonal and spatial dependence. The implicit aerosol treatment also tends to exclude days with high pollution, since aerosols are interpreted as effective clouds and the respective OMI pixels are often excluded by cloud screening; this is a potentially important sampling bias. Therefore an explicit treatment of aerosols

is important for space-based NO₂ retrievals and emission constraints. A comprehensive independent measurement network with sufficient spatial and temporal representativeness is needed to further evaluate the different satellite retrieval approaches.

1 Introduction

Tropospheric nitrogen oxides (NO_x = NO + NO₂) are important pollutants affecting ozone, aerosols, acid deposition, and climate. China has become the top emitter of NO_x due to its recent anthropogenic emission growth along with reductions in North America and Europe (Richter et al., 2005; van der A et al., 2008; Zhang et al., 2009; Lamsal et al., 2011; Castellanos and Boersma, 2012; Lin et al., 2014a). High NO_x pollution not only has significant consequences for China's domestic environment (Zhao et al., 2009; Lin et al., 2010a; Zhang et al., 2012), but it has also raised concerns regarding long-range pollution transport to downstream regions (Lin et al., 2008, 2014a; Cooper et al., 2010; Zhang et al., 2014; Jiang et al., 2015).

Vertical column densities (VCDs) of tropospheric nitrogen dioxide (NO₂) retrieved from the Ozone Monitoring Instrument (OMI) have been used extensively to study Chinese NO_x pollution (Stavrakou et al., 2008; Zhao and Wang, 2009; Lin et al., 2010b; Mijling et al., 2013; Miyazaki and Eskes, 2013). The high spatiotemporal coverage of OMI is superior to ground-based in situ measurements. However, NO₂ retrievals from OMI and other space-borne instruments are subject to large systematic and random errors due to uncertainties in the conversion process from radiance to VCDs (Boersma et al., 2011; Bucsela et al., 2013). In particular, current NO₂ algorithms take an implicit approach to accounting for aerosol optical effects, with no explicit specification of aerosols in the retrievals of both NO₂ VCDs and ancillary cloud parameters. The rationales for this approach are that (1) aerosols affect the retrieval of cloud parameters, so that the retrieved cloud parameters are “effective” and implicitly contain certain aerosol information, and (2) these effective cloud parameters at least partly describe the effect of aerosols on NO₂ air mass factors (Boersma et al., 2004, 2011). This implicit treatment is supported by the good spatial correlation (0.66) observed between coincident MODIS aerosol optical thickness values (mostly due to scattering) and O₂–O₂ effective cloud fractions over the eastern United States (Boersma et al., 2011).

Our previous study (Lin et al., 2014b) for several locations in the North China Plain (NCP) has shown large changes in retrieved NO₂ VCDs when moving from an implicit to an explicit treatment of aerosols. In particular, NO₂ VCDs are reduced by 14 % on average but are changed by (–90)–(+70) % for individual pixels when aerosol optical depth (AOD) exceeds 0.8. In addition, current NO₂ retrievals are

often based on monthly climatological surface albedo data, not accounting for the angular dependence of surface reflectance and its interannual variability; the corresponding effect on retrieved NO₂ has been estimated at 0–20 % for Europe (Zhou et al., 2010) and the NCP (Lin et al., 2014b) on average. Despite the complex terrains (Fig. 1a), complex land use types (Fig. 1b), and high aerosol loadings (Xin et al., 2007; Che et al., 2009; Wang et al., 2011) over China, the effects of aerosol and surface reflectance treatments are largely unknown.

This study extends our previous work (for a few locations; Lin et al., 2014b) to introduce an improved pixel-specific level-2 retrieval of tropospheric NO₂ VCDs over China (80–130° E, 20–53° N), Peking University OMI NO₂ (POMINO). Using a parallelized LIDORT-driven AMFv6 package (Palmer et al., 2001; Martin et al., 2003; Lin et al., 2014b), we explicitly account for aerosol optical effects, surface reflectance anisotropy, and their spatiotemporal variability. We then evaluate the individual and combined effects of an implicit aerosol treatment and changes in surface reflectance characteristics. In particular, we show large seasonal and spatial dependence of the effects of aerosol and/or surface reflectance treatments. We further illustrate the influences on subsequent NO_x emission constraints, a popular application of OMI data. Our POMINO data are available for 2004–2013 and will be updated to more recent times. Results for 2012 are presented here, by aggregating level-2 data into monthly mean values on a 0.25° long. × 0.25° lat. grid. Various provinces and regions are defined in Fig. 1a to facilitate the present analysis.

Section 2 presents our POMINO retrieval approach. Section 3 analyzes the POMINO NO₂ VCDs and the effects of various treatments of aerosols and surface reflectance. Section 4 further shows the effects on subsequent NO_x emission constraints. Section 5 concludes the present study with a further discussion on the applicability of our POMINO approach for a fast global retrieval from future fine-resolution satellite instruments.

2 Methodology

2.1 General process to retrieve tropospheric NO₂ VCDs

The OMI is a nadir-viewing imaging spectrometer onboard the polar-orbiting Sun-synchronous EOS Aura satellite with an Equator crossing time at 13:45 (Levelt et al., 2006). For each of the 14 or 15 orbits per day, the sensor measures UV-visible radiation reflected by the Earth from 60 across-track pixels with a swath of 2600 km. The pixel size is small and varies with the viewing angle (from 13 km × 24 km at nadir to 25 km × 150 km at the swath edge). The OMI pixels cover the globe on a daily basis, but the coverage of valid data is reduced by cloud and snow/ice contamination and

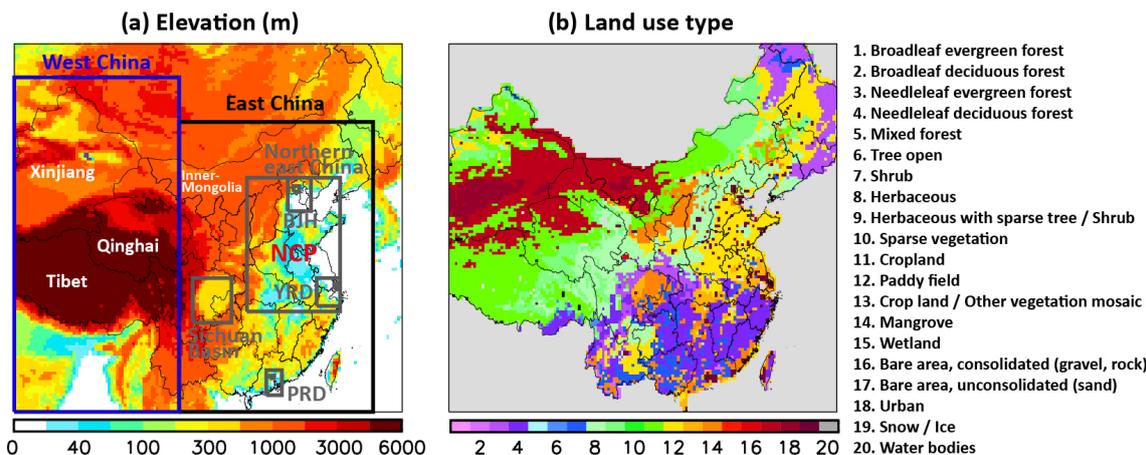


Figure 1. (a) USGS GMTED2010 surface elevation and (b) ISCGM GLCNMO land use types mapped to a 0.25° long. \times 0.25° lat. grid. Provincial boundaries of China are shown. Urban land use type is highlighted in (b), such that a grid cell is designated as “urban” if at least 5 % of its land is covered by urban areas. Also indicated in (a) are provinces (in white) and regions (in various colors) mentioned in the text. Eastern China: $101.25\text{--}126.25^\circ$ E, $20\text{--}46^\circ$ N; western China: $80\text{--}101.25^\circ$ E, $20\text{--}50^\circ$ N; northern eastern China: $110\text{--}122^\circ$ E, $29\text{--}41^\circ$ N; Sichuan Basin: $103\text{--}108^\circ$ E, $28\text{--}32^\circ$ N; Beijing–Tianjin–Hebei (BTH): $115.25\text{--}118.25^\circ$ E, $38\text{--}41^\circ$ N; urban Beijing: $116\text{--}116.75^\circ$ E, $39.75\text{--}40.25^\circ$ N; Yangtze River delta (YRD): $119\text{--}122^\circ$ E, $29.5\text{--}32^\circ$ N; urban Shanghai: $121.25\text{--}121.75^\circ$ E, $31\text{--}31.5^\circ$ N; Pearl River delta (PRD): $112.5\text{--}114.5^\circ$ E, $21.5\text{--}23.75^\circ$ N; and urban Guangzhou: $113\text{--}113.5^\circ$ E, $23\text{--}23.25^\circ$ N. Urban Beijing, urban Shanghai, and urban Guangzhou are inside BTH, YRD and PRD, respectively. The North China Plain (NCP) indicated in red represents the low-elevation (<300 m) areas of northern eastern China.

by recent row anomaly issues (especially since 2009). Row anomaly affects the quality of the level 1B radiance data for some viewing directions of OMI (<http://www.knmi.nl/omi/research/product/rowanomaly-background.php>).

Retrieval of tropospheric NO₂ VCDs from satellites normally undergoes a three-step process (Boersma et al., 2011). The first step derives slant column densities (SCDs) from satellite radiance data, and the second step separates the contribution of the tropospheric from the stratospheric part of the SCD. The final step involves an air mass factor (AMF) calculation to derive tropospheric VCDs (i.e., $\text{VCD} = \text{SCD} / \text{AMF}$). The AMF calculation is affected by surface reflectance, aerosol optical effects, cloud fraction (CF), cloud top pressure (CP), and atmospheric profiles of pressure, temperature and NO₂ (Zhou et al., 2010; Boersma et al., 2011; Bucsela et al., 2013; Lin et al., 2014b). Accurate knowledge of these parameters is an important prerequisite for the NO₂ retrieval. Over polluted regions like many parts of China, the AMF calculation is the dominant error source of retrieved tropospheric NO₂ data. Hereafter, SCDs, VCDs and AMFs are referred to as their tropospheric portions.

Before the main NO₂ retrieval itself, CF and CP are normally retrieved with the O₂–O₂ approach (Acarreta et al., 2004). Here, clouds are treated simply as an isotropically reflecting surface at a certain level (CP) with a Lambertian albedo of 0.8. As such, cloud optical properties are constrained by the two parameters CF and CP that can be retrieved from OMI data. Although O₂–O₂ cloud parameters are retrieved in a consistent manner with the latest DOMINO

v2 NO₂ retrievals (same surface albedo assumption, same radiative transfer model, same cloud model; Boersma et al., 2011) to ensure that the effective radiative properties of the scene are consistent between the cloud and NO₂ retrievals, some inconsistencies have recently come to light with respect to different atmospheric pressure and temperature profiles and terrain heights (Maasackers, 2013; Lin et al., 2014b).

2.2 Our POMINO retrieval approach

In this paper, our POMINO algorithm, referred to as case REF, largely follows the method described by Lin et al. (2014b) with a few modifications. Here we present a brief summary of the algorithm, placing emphasis on the latest modifications. The reader is referred to Lin et al. (2014b) for a detailed description. Our retrieval is focused on the derivation of tropospheric AMFs to calculate tropospheric VCDs, taking the tropospheric SCD data (Dirksen et al., 2011) from DOMINO v2 (Boersma et al., 2011). We adjust the calculated layer AMFs to correct for the temperature dependence of the NO₂ absorption cross section that is not accounted for in the SCD data (Boersma et al., 2004). Following Lin et al. (2014b), we re-retrieve the prerequisite (O₂–O₂-based) cloud properties by using a set of parameters consistent with those in the retrieval of NO₂. Our cloud retrieval is focused on AMF calculations, starting with the O₂–O₂ SCDs from the official cloud product OMCLDO2 v3 (Acarreta et al., 2004). Table 1 summarizes our retrieval approach and parameters. Figure 2 briefly illustrates the retrieval procedure.

Table 1. Key tools and parameters used in POMINO and associated cloud retrievals.

	POMINO NO ₂ retrieval	Cloud (O ₂ –O ₂)
AMF package	AMFv6; OpenMP parallelization	Same
RTM	LIDORT v3.6 (un-polarized, curved atmosphere); OpenMP parallelization	Same
Calculation for individual pixels	Pixel-specific radiative transfer modeling; no look-up table	Same
Surface reflectance	Land and turbid coastal ocean: BRDF at 440 nm, MCD43C2 Collection 5 (0.05°); open ocean: OMLER v3 albedo	Same
Surface pressure	GEOS-5 ^a ; adjusted by elevation ^b	Same
Cloud fraction and cloud pressure	Derived here	–
Aerosol optical parameters	GEOS-Chem v9-02 ^a ; at 438 nm; model AOD is adjusted by MODIS Aqua AOD ^c	Similar but at 475 nm ^c
Vertical profile of NO ₂	GEOS-Chem v9-02 ^{a,d}	–
Vertical profiles of pressure and temperature	GEOS-5 ^{a,d}	Same

^a Resolution: 0.667° long. × 0.5° lat. horizontally with 47 layers vertically and ~ 10 layers below 1.5 km. ^b Elevation information is from the GMTED2010 data set at 30 arcsec (http://topotools.cr.usgs.gov/GMTED_viewer/). ^c See Sect. 2.3 for details. ^d The pressure levels are re-calculated according to the elevation-adjusted surface pressure, while the volume mixing ratios of NO₂ are not changed in individual layers (Zhou et al., 2009).

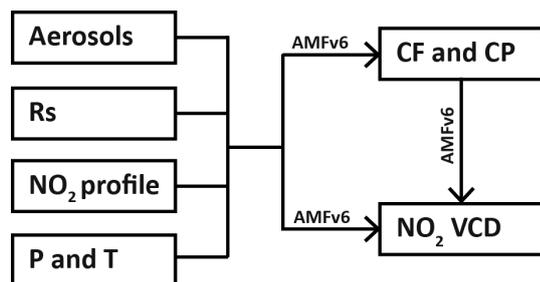


Figure 2. A diagram to illustrate our reference retrieval (case REF, representing our POMINO product). Here Rs is surface reflectance, P is air pressure, and T is air temperature. See Sect. 2.2 and Table 1 for detailed descriptions of our retrieval approach and parameters.

We calculated the AMFs for O₂–O₂ (to derive CF and CP; at 475 nm) and NO₂ (at 438 nm) by using a newly improved AMFv6 package (Palmer et al., 2001; Martin et al., 2003; Lin et al., 2014b) to coordinate the retrieval process. The AMFv6 code makes use of the LIDORT v3.6 radiative transfer model (RTM) (Spurr, 2008). Explicit radiative transfer is calculated pixel by pixel. Both the AMFv6 and LIDORT v3.6 codes have been parallelized via OpenMP. The parallelization has little overhead, in that, by using 16 computational cores in parallel, the speed-up is close to a factor of 16 relative to the single-core performance. With this speedup, a pixel-specific radiative transfer calculation becomes feasible for a large-scale (large domain, long time) retrieval, as in the present study. The traditional use of a look-up table to interpolate the AMFs is then no longer needed.

We used data for vertical profiles of NO₂, pressure and temperature on a relatively high-resolution grid (0.667° long. × 0.5° lat.). The NO₂ data were simulated by the GEOS-Chem chemical transport model (CTM), and pressure and temperature data were taken from the GEOS-5 assimilated meteorological fields that were used to drive GEOS-Chem simulations. GEOS-Chem has been shown to capture vertical profiles of NO₂ and ozone over the United States from aircraft measurements (Lin and McElroy, 2010). Appendix A summarizes the CTM simulations. As we retrieve clouds and NO₂ pixel by pixel, model information at the grid cell covering the pixel center is used. Although the size of our model grid cell is larger than the size of an OMI pixel, our model grid cell size is much smaller than that used in other OMI products [3° long. × 2° lat. for DOMINO (Boersma et al., 2011) and 2.5° long. × 2° lat. for OMNO2 (Bucsela et al., 2013)]. In addition, we adjust the pressure profile for each pixel based on the difference between pixel-specific surface elevation and grid cell average elevation (Zhou et al., 2009; Lin et al., 2012, 2014b). The meteorological and particularly NO₂ profiles are subject to errors (Boersma et al., 2011; Lin et al., 2014b). Further research is needed to evaluate these profiles using available measurements over China.

Our retrieval explicitly accounts for the effects of spatiotemporally varying aerosols and surface reflectance anisotropy on radiation. These two factors have proved relevant for the NO₂ retrieval (Zhou et al., 2010; Lin et al., 2014b; Noguchi et al., 2014). Detailed information is presented in Sects. 2.4 and 2.5.

2.3 Pixel selection, pixel-to-grid conversion, and sensitivity retrievals

For our present analysis, an OMI pixel is selected only when the ground is free from snow and ice, and when the cloud radiance fraction (CRF) does not exceed 50 % (Boersma et al., 2011; Lin et al., 2014b). Pixels with row-anomaly contamination are discarded. Data from valid pixels are then converted to monthly mean values on a 0.25° long. × 0.25° lat. grid through an area-weighted interpolation process. This process has been applied to NO₂ and all associated parameters. Section 3.5 discusses the number of days per month with valid pixels on the gridded map.

In addition to our POMINO retrieval (case REF), three other retrievals (cases SRF, AER, and S_A) were performed to evaluate the sensitivity of retrieved NO₂ VCDs and associated cloud parameters to changes in aerosols and surface reflectance. These additional retrievals, together with the standard DOMINO v2 product (referred to as case DOM) (Boersma et al., 2011), were compared with case REF. These tests are summarized in Table 2.

There are notable differences in the representation of CRF between POMINO and DOMINO. For POMINO, the CRF represents the fraction of the TOA radiance caused by clouds alone (in the context of additional contributions from the surface and aerosols). For DOMINO, however, the CRF applies to the fraction of TOA radiance caused by both clouds and aerosols, with surface reflectance represented by a geometry-independent surface albedo.

Different retrieval approaches lead to distinctive CRF values, which in turn has consequences for the selection of valid data (Lin et al., 2014b) (see discussions in Sect. 3.5). In Sects. 2 and 3, the pixels designated as “valid” by case REF are selected for analysis, regardless of their validity status in other retrievals. This choice ensures that the same set of pixels is evaluated for all retrieval methods. For the emission constraint study in Sect. 4, different sets of valid pixels specific to the individual retrieval approaches are also analyzed, in addition to the set determined by case REF.

2.4 Surface reflectance in POMINO (case REF)

Accurate knowledge of surface reflectance is of key importance for retrievals of NO₂ and ancillary cloud parameters. Surface reflectance depends both on the ground conditions and on the state of the overlying atmosphere (the latter determining the relative amounts of diffuse vs. direct incident radiance) (Lucht et al., 2000). Due to inhomogeneity in surface conditions, the amount of reflected radiance relative to a given amount of incident radiance depends on the incoming and outgoing angles. The degree of angular dependence is determined by surface roughness and vegetation characteristics (type, leaf area, and geometric shape). The angular dependence is most prominent for direct incident radiance; angular effects on the reflected radiation field are largely can-

celled out for diffuse (isotropic) incident radiance (Lucht et al., 2000). Normally the angular dependence is not accounted for in the NO₂ retrievals, with a few exceptions, e.g., Zhou et al. (2010) for Europe, Noguchi et al. (2014) for Tokyo, and Lin et al. (2014b) for several locations in China.

Case REF explicitly accounts for the angular dependence of surface reflectance. It adopts the MODIS BRDF product as an approximate realization of the complex dependence of surface reflectance on radiation geometry (Lucht et al., 2000). This product is based on BRDF models using a linear combination of three near-independent reflecting kernels: isotropic (no angular dependence), volumetric (related to leaf area), and geometric (related to vegetation shape). Each kernel contribution is regulated by a parameter that varies with time and space. We used the snow- and ice-free MCD43C2 Collection 5 data set (Lucht et al., 2000) that provides the kernel parameters over land and turbid coastal ocean at 440 nm as 16-day average values on a 0.05° long. × 0.05° lat. grid. Kernel parameters are updated every 8 days, accounting for the seasonal and interannual variability in BRDF. To cover missing values and reduce noise, the high-resolution data were passed through a spatiotemporal interpolation and smoothing procedure. The data were then mapped to each OMI pixel for subsequent AMF calculations. A detailed description of these customized procedures is presented in Lin et al. (2014b). Figure S1a–c in the Supplement presents the horizontal distributions of BRDF kernel parameters on the original 0.05° long. × 0.05° lat. grid for the time period of 25 June 2012–10 July 2012.

Over China, the angular dependence of surface reflectance is more important over parts of the west, southwest, southeast and northeast having complex terrains (Fig. 1a), vegetated lands (Fig. 1b), and relatively low aerosol loadings (Fig. 3, first row). Over the NCP and many other polluted regions of China, high aerosol loadings (Fig. 3, first row) mean that most incident radiation to the ground is diffuse, so that the angular dependence is reduced for the radiance reflected from the ground. Nonetheless, the seasonal and interannual variability in surface reflectance may still be important for China, especially considering rapid land use change due to urbanization, industrialization and agricultural activities (Liu et al., 2014).

Over the open oceans, there are no applicable BRDF data; therefore, we used the surface albedo data from the OMLER v3 product on a 0.5° long. × 0.5° lat. grid (Kleipool et al., 2008). This albedo data set is a 5-year (2005–2009) mean monthly climatology, an update of OMLER v1 adopted by DOMINO v2. The oceanic pixels are included for completeness in our product; exclusion or inclusion of these pixels has little effect on our present analysis. Figure S1d in the Supplement shows the OMLER v3 data in July.

Figure 4 (first row) presents the horizontal distribution of annual and seasonal bi-directional reflectance factor (BRF) values in 2012, as representative of MODIS BRDF over land and turbid coastal ocean and OMLER v3 albedo over the

Table 2. Retrievals of clouds and NO₂ with different approaches*.

	REF = POMINO	SRF	AER	S_A	DOM
Surface reflectance	MCD43C2 BRDF	OMLER v3 albedo	MCD43C2 BRDF	OMLER v3 albedo	OMLER v1 albedo
Aerosol treatment	Explicit	Explicit	Implicit	Implicit	Implicit

* Retrieval procedures and parameters not mentioned here are the same for cases REF, SRF, AER and S_A. Case DOM is adopted from DOMINO v2.

open ocean. The data have been sampled from valid OMI pixels and mapped to a 0.25° long. × 0.25° lat. grid, as described in Sect. 2.3. The BRDF here is the ratio of reflected radiance to the π -divided direct incident irradiance (along the geometric light path from the Sun to the ground and then to the OMI) (Schaeferman-Strub et al., 2006). Especially for incident radiation with strong direct and weak diffuse contributions, the value of BRDF is critical for the total amount of radiation received by OMI. The BRDF data in Fig. 4 (first row) imply that the spatial and temporal variations in solar zenith angle are taken into account, in addition to changes in the ground characteristics. The choice of BRDF for presentation purposes follows Zhou et al. (2010) and Lin et al. (2014b).

Figure 4 (first row) shows that in all seasons, the BRDF reaches maximum values of 0.1–0.3 over western and northwestern China with desert or bare land. The minimum values at 0.02–0.04 occur in parts of eastern China where there is stronger absorption by vegetation (Fig. 1b). Especially over eastern China, the relatively high (low) BRDF values are often coincident with high (low) AOD (comparing Figs. 3 and 4, first rows), likely indicating the presence of aerosol contamination in the BRDF data. The BRDF data exhibit significant seasonal variation. Over eastern China, the BRDF often reaches maximum values in summer and minima in fall. For example, over much of the NCP, the BRDF varies from 0.06–0.08 in summer to 0.04–0.06 in fall. Over the west, the summer season has the lowest BRDF values, likely a result of a lower solar zenith angle.

2.5 Aerosol optical properties in POMINO (case REF)

POMINO explicitly accounts for the optical effects of aerosols, given the high aerosol loadings over China (Fig. 3, first row). Lin et al. (2014b) has a detailed description of the implementation of aerosol optical properties in the retrieval process. Here we emphasize the modifications to POMINO needed to facilitate a large-scale retrieval (i.e., for a large domain in all seasons, as compared to several spot locations investigated by Lin et al., 2014b).

Day-to-day varying aerosol optical properties (AOD, single scattering albedo (SSA), phase functions, and vertical profiles) are taken from the GEOS-Chem v9-02 simulations on a 0.667° long. × 0.5° lat. grid. The model is updated from an earlier version (v8-03-02) used by Lin et al. (2014b). See Appendix A for model descriptions. GEOS-Chem simulates various aerosol types, including secondary inorganic aerosols

(sulfates, nitrates, and ammoniums), organic aerosols, black carbon, dust and sea salts. For a given OMI pixel, aerosol data at the grid cell covering the pixel center are used during the retrieval process. Aerosol optical properties at two wavelengths are implemented to retrieve NO₂ (438 nm) and clouds (475 nm), respectively.

To constrain the AOD, CTM-modeled AOD at 550 nm is adjusted to match MODIS Aqua data on a monthly basis (Appendix B). The AOD adjustment is then carried over to other wavelengths (438 and 475 nm) based on species-specific size distributions, refractive indices and hygroscopic growth rates as assumed in GEOS-Chem. The same procedure was used in Lin et al. (2014b).

Figure 3 (first row) presents the horizontal distribution of AOD at 550 nm used in POMINO. Both annual and seasonal mean data are shown. High AOD values are apparent over the NCP (the annual mean is 0.8–1.0), the Sichuan Basin (0.8–1.0), and parts of southern China (0.6–0.8) due to significant anthropogenic sources. The high AOD values over the Sichuan Basin are also a result of a long-lasting stagnant atmosphere. Large AOD values are also present over the western deserts, especially in spring (0.9–1.2), which has the highest dust emissions. Over eastern China, AOD values are higher in spring and summer than in fall and winter. These spatial and temporal patterns are generally consistent with previous findings (Xia et al., 2007; Xin et al., 2007; Wang et al., 2011; Lin et al., 2014c).

Figure 3 (second row) shows the SSA at 550 nm. The SSA is largest over western and northwestern China, where there are few black carbon sources. Over the west, the SSA varies between 0.92 and 0.98 in all non-winter seasons. In winter, the SSA is reduced to 0.90–0.92 over large parts of Xinjiang. Over eastern China, the SSA experiences even larger seasonal and spatial variability, from ~0.80 over parts of the NCP in winter to 0.94–0.96 over most of eastern China in summer. The seasonality of SSA is mostly a consequence of black carbon emissions reaching their maximum values in winter and minimum amounts in summer (see Appendix A for the implementation in GEOS-Chem).

Figure 3 (second row, filled circles) shows SSA values for 2005 estimated by Lee et al. (2007) from MODIS top-of-atmosphere radiance and ground AOD networks. Their estimates are only for situations with AOD > 0.4, and have a root mean square error of 0.03. Differences between our SSA values and Lee et al. are highly season- and location-dependent. On a seasonal basis, our SSA values, sampled at grid cells

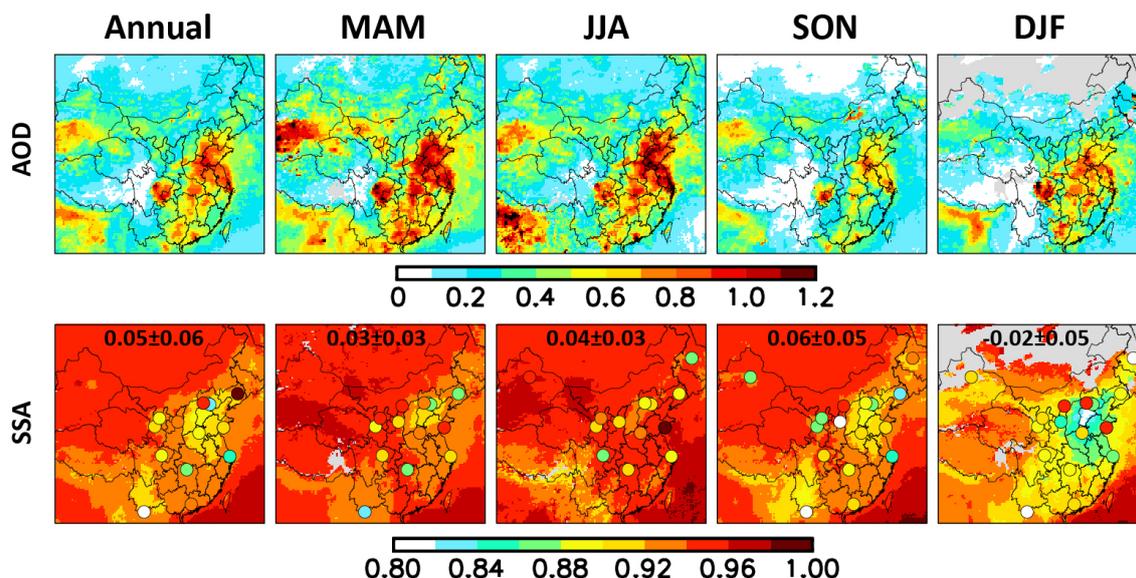


Figure 3. AOD and SSA at annual and seasonal scales. Provincial boundaries of China are shown. Data are sampled from valid pixels of case REF. AOD values exceeding a value of 1.2 are shown in black. Missing values are shown in grey. Filled circles in the second row indicate the SSA estimates for 2005 with AOD > 0.4 by Lee et al. (2007). The embedded numbers are our SSA values minus Lee et al. (± 1 standard deviation); our data are sampled at grid cells covering their sites.

covering their sites, are most consistent with Lee et al. in winter (mean difference across China is -0.02 ± 0.05), followed by spring, summer and fall. Note that these comparisons are qualitative, given the inconsistency in data sampling.

Several limitations constrain our ability to improve aerosol modeling. Model aerosol optical properties (AOD, SSA, phase functions) and vertical profiles are subject to errors (Drury et al., 2010; Ford and Heald, 2012; van Donkelaar et al., 2013). We used MODIS AOD data to constrain CTM-derived AOD, even though MODIS data are not free of errors (Wang et al., 2007; Wang et al., 2010; Hyer et al., 2011). No adequate observations are available to constrain other aerosol optical parameters at a regional scale with high spatial and temporal resolutions. Observation-based estimates of SSA are essentially lacking at the scale considered here, and the few results in the literature contain large uncertainties (± 0.03) (Lee et al., 2007). Although the CALIOP instrument provides information of aerosol vertical profiles (Winker et al., 2009), the CALIOP profiles are limited by their spatiotemporal coverage and data quality (especially near the ground) (Ford and Heald, 2012; van Donkelaar et al., 2013). Note that since the same vertical mixing and convection schemes were used to simulate aerosols and NO₂, the height of aerosols relative to NO₂ (relevant to our study) may be subject to smaller errors than the absolute height of aerosols. Future work is needed to better understand and constrain aerosol properties and evaluate how they affect the NO₂ retrieval.

Nevertheless, our present study, at the very least, reveals the importance of an explicit aerosol treatment for NO₂ and

associated cloud-parameter retrievals at a regional scale, especially given the lack of such an explicit treatment in current satellite products. In support of our work here, Lin et al. (2014b) showed that, by explicitly accounting for aerosols with just the AOD values constrained by observations, there is excellent correlation between retrieved NO₂ VCDs and independent MAX-DOAS data ($R^2 = 0.96$ in day-to-day variability across the few locations being studied). Section 3.3 further shows large changes in retrieved NO₂ VCDs from an explicit to an implicit treatment of aerosols, and Sect. 4 illustrates the consequences for subsequent NO_x emission constraint. Therefore, we expect that the explicit inclusion of aerosols will improve the NO₂ retrieval, especially if more comprehensive observations become available to constrain model aerosols.

3 OMI NO₂ retrievals and complex influences of aerosols and surface reflectance

3.1 General characteristics of POMINO NO₂ (case REF)

Figure 5 (first row) shows annual and seasonal mean NO₂ VCDs from POMINO (case REF) on a $0.25^\circ \text{ long.} \times 0.25^\circ \text{ lat.}$ grid. At this fine resolution, hotspots of NO₂ pollution across China are clearly visible. The formation of hotspots is also a result of the short lifetime of NO₂ (2–3 h to 1 day, depending on chemical activity). Large pollution covers much of the NCP, with annual mean NO₂ VCDs exceeding $15 \times 10^{15} \text{ cm}^{-2}$, due to emissions from both urban

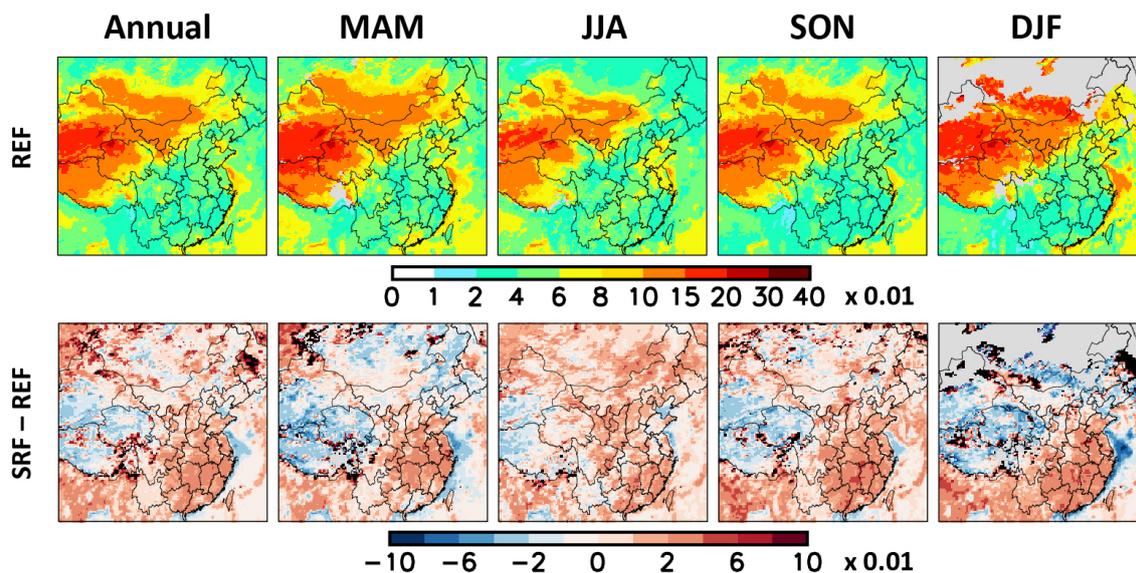


Figure 4. First row: surface reflectance in case REF = POMINO (MODIS BRF data) at annual and seasonal scales. Second row: surface reflectance in case SRF (OMLER v1 albedo) minus reflectance in case REF (MODIS BRF). Provincial boundaries of China are shown. Data are sampled from valid pixels of case REF. Values outside the upper (lower) bound of color intervals are shown in black (purple). Missing values are shown in grey. Color intervals are nonlinear to better present the data range; an interval without labeling represents the mean of the adjacent two intervals.

and regional sources. Annual mean NO₂ VCDs are below 10^{15} cm⁻² over much of western China due to a lack of anthropogenic influences. The spatial gradient of NO₂ pollution is greatest in summer due to its shortest lifetime, but the NCP still has a large inter-connected area of high NO₂. This regional-scale high pollution highlights the severity and extensiveness of China's environmental problems.

Figure 5 (first row) also shows a large seasonal variation in NO₂. Over eastern China, NO₂ VCDs reach maxima in winter and minima in summer. The maximum to minimum ratio is about 3.6 for all regions east of 101.25° E. The large seasonality in NO₂ VCDs mainly reflects the seasonality in the species' lifetime (Lin, 2012). Over most of western China with few anthropogenic emissions, NO₂ VCDs are largest in summer due to a peak in natural (lightning and soil) sources that overcompensates for the shortest lifetime. This seasonal pattern is most notable over Tibet and Qinghai. For western China (west of 101.25° E) as a whole, the ratio of summer peak to winter minimum is about 1.4. Over much of Xinjiang and Inner Mongolia, the growth in anthropogenic influences has meant a winter maximum and a summer minimum, reversing the seasonality typical for western China.

Figure 5 (fifth row) presents the difference in NO₂ VCDs between the cases DOM and REF, as a percentage fraction of REF. Case DOM is taken from DOMINO v2 (Boersma et al., 2011) and sampled from pixels valid in case REF, irrespective of whether these pixels are flagged as valid in the DOMINO v2 product. (Note that in Sect. 4, valid pixels from case REF and DOMINO v2 are both evaluated for the

derivation of NO_x emissions.) Figure 5 (fifth row, first panel) shows that at an annual scale, results from case DOM exceed those from REF by 0–60 % over central eastern China (consistent with POMINO NO₂ columns being ~45 % lower than DOMINO, as reported in Lin et al., 2014) and much of the west. Case DOM results are smaller than those from REF over parts of the south and north. Seasonal dependence is significant (Fig. 5, fifth row). Over the NCP, case DOM is greater than REF by 10–40 % in summer, while the signs of difference are location-dependent in winter. Over most of Tibet, case DOM is similar to REF in the fall season, but greatly exceeds REF (by up to 40 %) in other seasons. These differences reflect the dissimilar AMF approaches in the two retrievals. It is beyond the scope of this study to fully elucidate the differences between case REF (POMINO) and DOMINO v2. Instead, the following sections analyze the effects of surface reflectance and aerosols on retrieved NO₂ VCDs.

3.2 Effects of surface reflectance on NO₂ and cloud retrievals

The MODIS BRDF data account for spatial and temporal variability in surface reflectance as well as its angular dependence. Here we evaluate the sensitivity of retrieved NO₂ VCDs to surface reflectance, by adopting an alternate surface reflectance data set to repeat the retrieval process.

Case SRF adopts the OMI monthly climatological albedo data (OMLER v1, at 440 nm) from the DOMINO v2 product and re-derives cloud properties and NO₂ VCDs; other retrieval procedures, including aerosol treatments, are un-

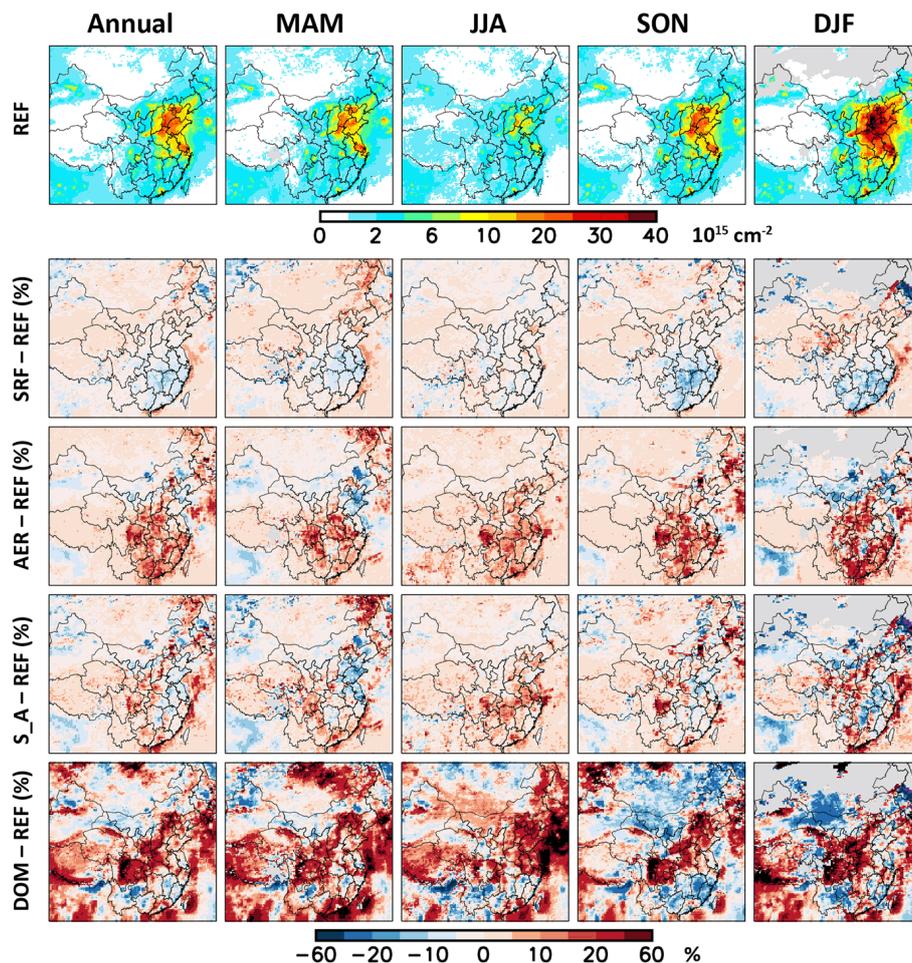


Figure 5. First row: tropospheric NO₂ VCDs retrieved from case REF at annual and seasonal scales. Second–fifth rows: changes in NO₂ VCDs from case REF to other cases as a percentage fraction of case REF. Provincial boundaries of China are shown. Data are sampled from valid pixels of case REF. Values outside the upper (lower) bound of color intervals are shown in black (purple). Missing values are shown in grey. Color intervals are nonlinear to better present the data range; an interval without labeling represents the mean of the adjacent two intervals.

changed. Unlike the MODIS BRDF data, the OMI albedo data are monthly climatology (October 2004–October 2007 average) with no interannual variability. The OMI albedo data set means isotropic reflectance with no angular dependence. The horizontal resolution of OMI albedo data is 0.5° long. \times 0.5° lat., compared to the high-resolution MODIS BRDF data (at 0.05° long. \times 0.05° lat.). The use of OMI albedo data tests the sensitivity of retrieved NO₂ to large changes in surface reflectance.

Figure 4 (second row) compares the OMI albedo with the MODIS BRF over China. The OMI albedo is normally within ± 0.05 of the MODIS BRF. Over most of eastern China, the OMI albedo exceeds the MODIS BRF in all seasons with a difference of 0.01–0.04. Over northeastern China, however, the OMI albedo is lower than the MODIS BRF in spring and winter. Over most of western China, the OMI albedo is smaller than the MODIS BRF with a dif-

ference of 0.01–0.06. The OMI albedo greatly exceeds the MODIS BRF at various locations in the west and north (by 0.10 or more).

The diagram in Fig. 6 illustrates how a change in surface reflectance affects the pre-NO₂ cloud retrieval. In the cloud-property O₂–O₂ algorithm, higher reflectance leads to lower effective CF, since fewer clouds are needed to reflect a given amount of radiation to the outer space. Effects of changing surface reflectance on CP are multifold. Higher reflectance means an enhanced AMF in the clear-sky portion of the OMI pixel (AMF_{cr}), which can be compensated for by a decrease in CP. In addition, the reduction in CF caused by enhanced surface reflectance has a secondary effect on CP. A decrease in CF may lead to a further reduction in CP if the AMF of O₂–O₂ in the cloudy portion of the OMI pixel, AMF_{cl}, is smaller than AMF_{cr} (this is a “screening” effect of clouds on radiation). This effect occurs in most situations where the

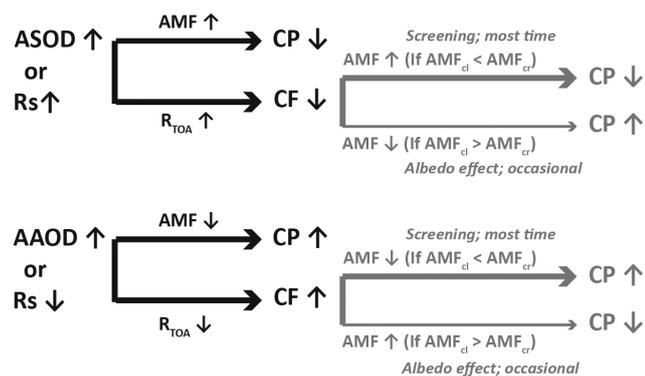


Figure 6. A diagram of how changes in surface reflectance (Rs), aerosol scattering (ASOD) and aerosol absorption (AAOD) affect the retrievals of CF and CP in the O₂–O₂ algorithm. Grey colors illustrate the indirect influence of CF changes on the CP; the influence depends on the “screening” or “albedo” effects of clouds on radiation. Thin grey lines indicate that the “albedo” effect of clouds is occasional. The figure is updated from Fig. 5 of Lin et al. (2014b).

cloud top is distant from the ground and there is low above-cloud O₂–O₂ concentration. Occasionally, a decrease in CF may result in an enhancement in CP, when the cloud top is close to the ground and thus the value of AMF_{cl} exceeds AMF_{cr} (an “albedo” effect of clouds on radiation).

Figure 7 contrasts case REF and SRF for the CF, cloud radiance fraction (CRF) and CP at an annual scale. A negative correlation is apparent between changes in CF (and CRF) from case REF to SRF and changes in surface reflectance. Over land, the CP tends to change in the opposite direction to the change in surface reflectance. Over much of the oceans away from Chinese coasts, however, the CP decreases with declining surface reflectance from case REF to SRF; this is due to the “albedo” effect of clouds (CP = 800–900 hPa). Overall, the CP changes from case REF to SRF with an opposite sign as surface reflectance for about 78 % of all grid cells and months. Changes are similar across different seasons for individual cloud parameters (Figs. S2–S4 in the Supplement, second rows).

As shown in Eq. (1) (the IPA or independent pixel approximation), changes in CRF affect the relative weights of NO₂ AMF_{cl} vs. AMF_{cr}, while changes in CP affect the absolute magnitude of NO₂ AMF_{cl}. In addition, an increase in surface reflectance results in an enhancement in NO₂ AMF_{cr}. These factors together determine the effects of surface reflectance on the NO₂ retrieval.

$$\text{AMF} = \text{AMF}_{cl} \cdot \text{CRF} + \text{AMF}_{cr} \cdot (1 - \text{CRF}) \quad (1)$$

$$\text{VCD} = \frac{1}{1/\text{VCD}_{cl} \cdot \text{CRF} + 1/\text{VCD}_{cr} \cdot (1 - \text{CRF})} \quad (2)$$

Figure 5 (second row) shows that the effects of surface reflectance on retrieved NO₂ VCDs are largely region- and season-dependent. Over most Chinese regions except central

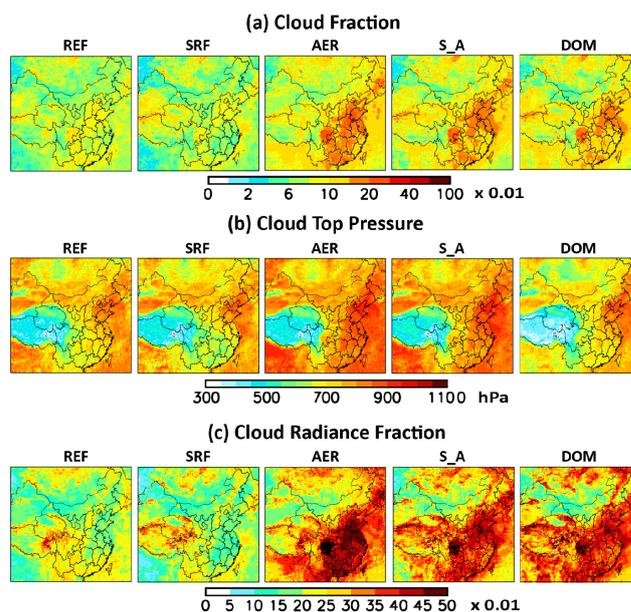


Figure 7. Annual mean (a) cloud fraction, (b) cloud top pressure and (c) cloud radiance fraction retrieved via different approaches. Provincial boundaries of China are shown. Data are sampled from valid pixels of case REF. Values outside the upper (lower) bound of color intervals are shown in black (purple). Missing values are shown in grey. Color intervals in (a) are nonlinear to better present the data range; an interval without labeling represents the mean of the adjacent two intervals.

and southeastern China, replacing the MODIS BRDF with OMI albedo tends to increase the retrieved NO₂ VCDs by 0–40 % (mostly less than 15 %). Over central and southeastern China, however, the use of OMI albedo reduces NO₂ VCDs by 0–20 % in spring, fall and winter, but with a slight enhancement in summer. This spatial and seasonal divergence in NO₂ changes reflects the complex influences of surface reflectance on retrieved cloud properties and NO₂ AMFs. The magnitude of NO₂ changes is comparable to previous studies (Zhou et al., 2010; Lin et al., 2014b).

Figure 8b presents the percentage changes in NO₂ VCDs from case REF to SRF as a function of AOD values (binned at intervals of 0.05) and changes in surface reflectance (i.e., OMI albedo minus MODIS BRDF, binned at intervals of 0.01). Here the percentage changes from all grid cells and months with respect to each bin of AOD and surface reflectance change are averaged; the frequency of data located in each bin is shown in Fig. 8a. Figure 8c and d also separate the effects of surface reflectance on NO₂ VCDs for the pixel clear-sky and cloudy portions. Here, a clear-sky VCD (VCD_{cr}) is derived as the SCD divided by AMF_{cr}, and a cloudy-sky VCD (VCD_{cl}) represents the SCD divided by AMF_{cl}; Eq. (2) shows the relation between VCD, VCD_{cr} and VCD_{cl}. Pixels with no clouds are excluded from calculations of AMF_{cl} and VCD_{cl}. The VCD_{cr} and VCD_{cl} for individual pixels

are aggregated to respective monthly mean values in the same way as was done for the total VCD.

Figure 8c shows that, when the OMI albedo is larger (smaller) than the MODIS BRDF, we get a reduction (enhancement) in VCD_{cr}. The relative changes in VCD_{cr} are within 40 % (mostly < 10 %), and a greater change in surface reflectance tends to result in a larger change in VCD_{cr}. In any given AOD bin, the magnitude of correlation between changes in surface reflectance and changes in VCD_{cr} exceeds 0.8 (embedded red line). For a given bin of surface reflectance, by comparison, there is no apparent dependence of VCD_{cr} changes on the amounts of AOD. The effects of surface reflectance on NO₂ VCDs are similar to the effects on VCD_{cr} (Fig. 8b) due to the low CRF values on average (Fig. 7c, second column).

Figure 8d shows complex effects of surface reflectance perturbations on VCD_{cl}. This is because the CP does not always change in the same direction as surface reflectance (see discussions above). When the OMI albedo is within ± 0.05 of the MODIS BRDF, the VCD_{cl} tends to increase from cases REF to SRF. There is greater scatter in VCD_{cl} changes when the changes in surface reflectance are greater. In addition, VCD_{cl} undergoes a much greater change in magnitude than VCD_{cr}, reflecting the strong sensitivity of AMF_{cl} to the CP.

3.3 Influences of implicit aerosol treatment on NO₂ and cloud retrievals

Case AER tests the effects of aerosols on NO₂ and cloud retrievals by setting AOD to zero during the retrieval process, thus mimicking the traditional treatment (Boersma et al., 2011; Bucseila et al., 2013). This procedure leads to changes in CF and CP that implicitly affect the subsequent NO₂ retrieval. If this implicit treatment results in the same NO₂ VCDs as case REF, then an explicit treatment of aerosols is no longer strictly desirable, as it is more expensive computationally.

Figure 7a and c show that at an annual scale, the exclusion of aerosols results in significant enhancements in CF and CRF. This is because aerosols reflect solar radiation to space, and exclusion of aerosols is compensated for by an increase in effective CF and CRF (see Fig. 6) (Boersma et al., 2011; Lin et al., 2014b; Castellanos et al., 2015). Over China, the CF is enhanced from 0.04–0.15 in case REF to 0.06–0.30 in case AER. The increase is greatest over parts of eastern China, with large AOD values in case REF. The CRF in case AER is greater than 0.35 over most of eastern China, compared to the values of 0.15–0.30 in case REF. Over the Sichuan Basin, the annual mean CRF exceeds 0.50 in case AER, more than doubling the CRF value (about 0.25) in case REF. Overall, the correlation between the amounts of AOD neglected in case AER and the level of CF (CRF) augmentation reaches a high value of 0.75 (0.82). Analyses of individual seasons also show large enhancements in CF and CRF as a result of neglecting aerosols (Figs. S2 and S4 in the Supple-

ment, third rows). Correlation between effective CF and AOT has been found for the eastern United States (0.66) (Boersma et al., 2011) and South America (Castellanos et al., 2015).

Figure 7b shows that excluding aerosols leads to an increase in CP from case REF to AER, in order to compensate for an otherwise reduction in the O₂–O₂ AMF. Over eastern China, the CP is increased from 700–900 hPa in case REF to 750–950 hPa in case AER. The CP enhancement is smaller over western China, and is smallest (0–20 hPa) over the Tibetan Plateau due to the lowest aerosol loadings. However, the correlation between the amounts of AOD neglected in case AER and the amounts of CP increase is only 0.19, reflecting the complex effect of aerosols on the CP retrieval (Lin et al., 2014b). The CP enhancements are apparent in all seasons (Fig. S3 in the Supplement, third row).

Figure 5 (third row) shows the horizontal distribution of percentage NO₂ changes from cases REF to AER. In all seasons, case AER is larger than REF by 0–40 % over most of China. The overestimate is most obvious in central and southern China, especially in winter (by 15–40 %). By comparison, case AER leads to lower NO₂ VCDs by 0–20 % over parts of the NCP in spring, many places in the north in winter, and parts of the west in non-summer seasons. At an annual scale, case AER leads to larger NO₂ VCDs than case REF by 0–40 % over most regions. The magnitudes of NO₂ changes are weakly correlated with the AOD or SSA (comparing Fig. 5, third row with Fig. 3). The degree of divergence in NO₂ changes is consistent with that found in Lin et al. (2014b).

The red solid lines in Fig. 9 present the AER to REF ratio for NO₂ VCDs on a monthly basis. Several representative regions of China are considered, including eastern China, western China, northern eastern China, the Sichuan Basin, Beijing–Tianjin–Hebei (BTH), the Yangtze River delta (YRD), the Pearl River delta (PRD), and urban areas of Beijing, Shanghai and Guangzhou. These regions are defined in Fig. 1a. Figure 9 shows that for NO₂ VCDs averaged over a large region such as eastern or western China, the AER to REF ratio is close to 1 (0.9–1.1) in all months. For smaller regions, the deviation in the AER/REF ratio increases dramatically. The ratios vary between 0.8 and 1.4 across the 12 months for BTH, YRD and PRD, and between 0.8 and 1.6 for the urban areas of Beijing, Shanghai and Guangzhou. Whether case AER leads to larger or smaller NO₂ VCDs than case REF depends strongly on location and season. For urban Beijing, the maximum AER/REF ratio over the 12-month period is about twice as much as the minimum AER/REF ratio. Case AER has larger NO₂ VCDs than case REF in most months over the Sichuan Basin and the PRD, while the AER/REF ratios are more seasonally variable in other regions. These regional and seasonal features call for comprehensive independent measurements to validate satellite retrievals.

Figure 10b shows the percentage changes in NO₂ VCDs from case REF to case AER as a function of AOD and SSA.

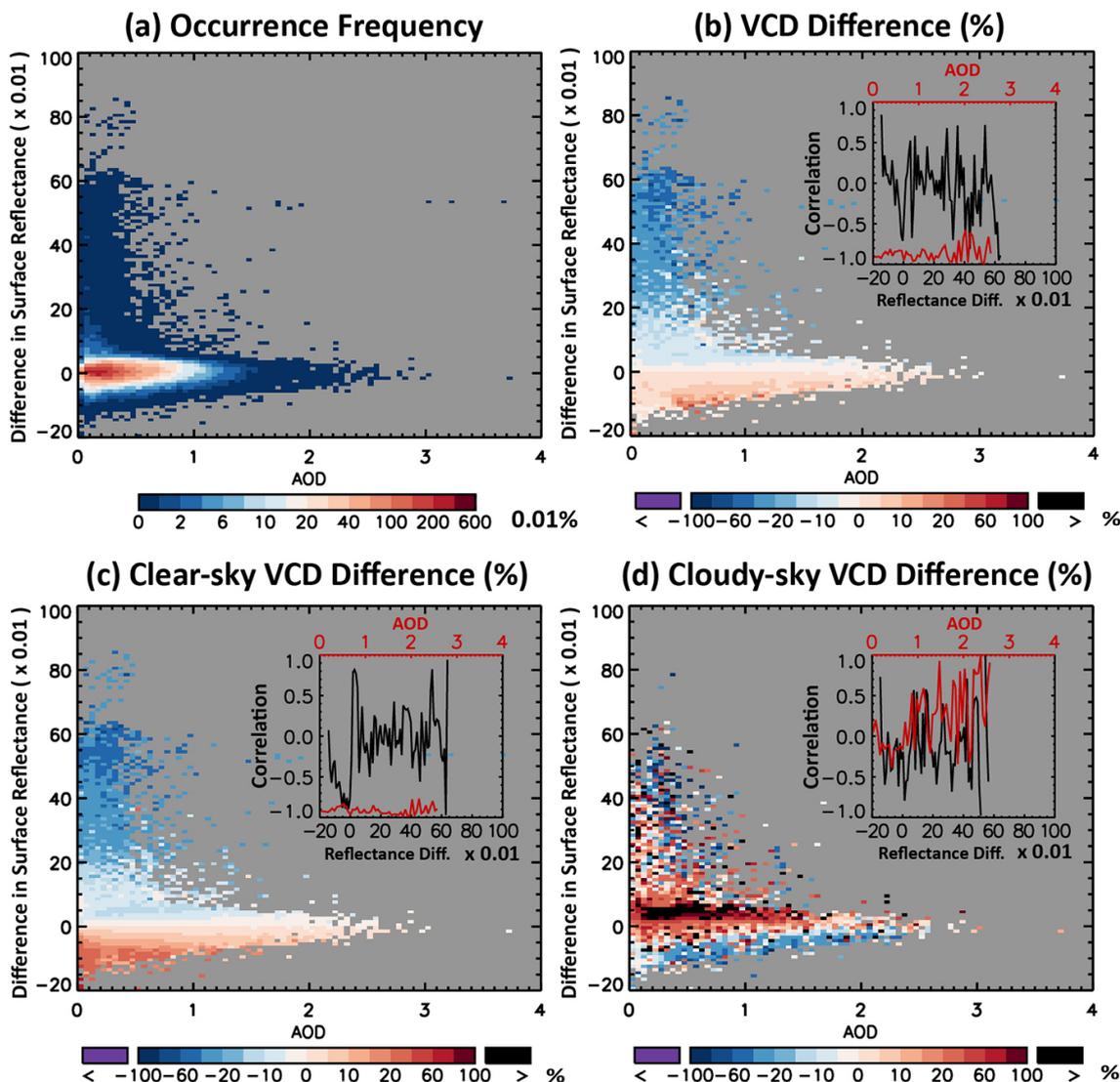


Figure 8. (a) Frequency of occurrence from all months and grid cells for each bin of AOD (x axis, bin size 0.05) and surface reflectance difference (OMI albedo in case SRF minus MODIS BRDF in case REF, y axis, bin size 0.01). Data are sampled from valid pixels of case REF. (b) Percentage changes in NO₂ VCDs from case REF to SRF averaged over all data in each bin of AOD and surface reflectance difference; also embedded are the correlations between percentage VCD changes and AOD values as a function of surface reflectance differences (black line) and between VCD changes and surface reflectance differences as a function of AOD values (red line). Similar panels are drawn for (c) clear-sky VCDs and (d) cloudy-sky VCDs.

Here the percentage changes from all grid cells and months with respect to each bin of AOD (bin size = 0.05) and SSA (bin size = 0.01) are averaged; the amount of data in each bin is shown in Fig. 10a. Whether a larger AOD value corresponds to a greater enhancement in NO₂ VCD from case REF to case AER depends in a complex manner on the SSA. This effect is also found by Castellanos et al. (2015) for biomass burning aerosols over South America. The largest increase (by 40–80 %) from case REF to case AER occurs with high AOD of 1–2 and low SSA of ~ 0.90 . The dependence of NO₂ changes on SSA is weak, with positive cor-

relations when AOD is below 1.5 and negative correlations when AOD exceeds 1.5.

Figure 10c presents the percentage changes in NO₂ VCDcr as a function of AOD and SSA. With a fixed AOD value, lower SSA tends to result in a smaller increase or a large decrease in VCDcr from case REF to AER; for AOD of 0.2–1.3, the correlation between VCDcr changes and SSA values exceeds 0.7. This is because stronger absorption diminishes the radiation that could otherwise be absorbed by NO₂, and leads to a consequent reduction in AMFcr. The dependence of VCDcr changes on AOD is highly SSA-sensitive. In many situations, the correlation between NO₂ changes and AOD

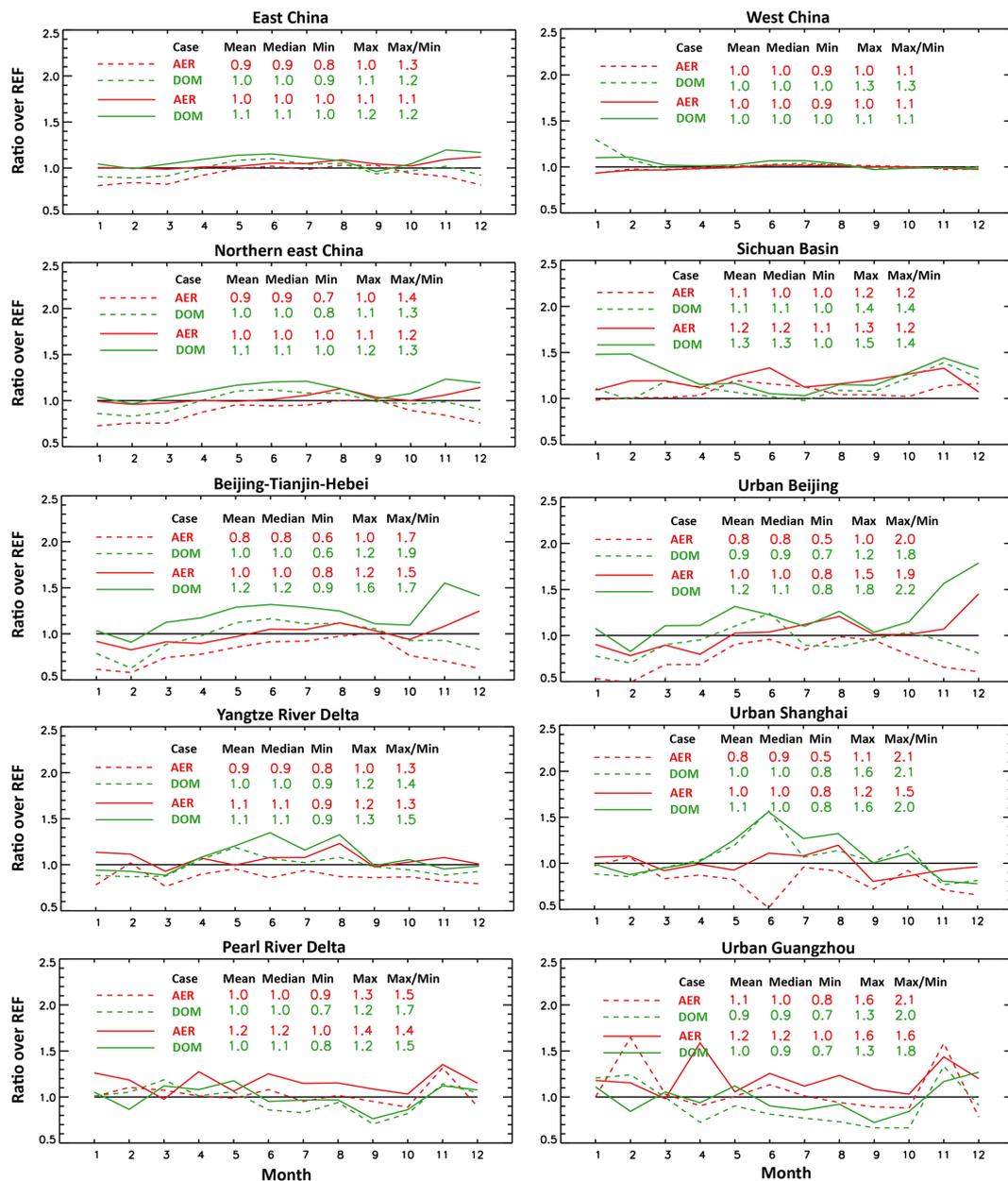


Figure 9. Ratios of cases AER and DOM to case REF for regional mean NO₂ VCDs in each month of 2012. Regions are defined in Fig. 1a. For cases AER and DOM, data can be sampled from pixels valid in respective retrieval approaches (dotted lines) or valid in case REF (solid lines). Also shown are some statistical quantities.

or SSA values is weak (within ± 0.5), owing to the indirect effects of aerosol vertical profiles and other factors. These features reflect the complex effects of aerosol extinction on the radiation absorbed by NO₂ under cloud-free conditions, as has been found in GEOS-Chem simulations (Lin et al., 2012). This contrasts with the effect of changing surface reflectance on NO₂ VCD_{cr} such that an increase in surface reflectance reduces the amount of VCD_{cr} (Fig. 8c).

Figure 10d shows that the effect of excluding aerosols on VCD_{cl} differs significantly from the effect on VCD_{cr}. Case

AER leads to lower VCD_{cl} values by 0–60% for a wide range of AOD and SSA, mostly because the increase in CP leads to a consequent enhancement of AMF_{cl}.

We further elucidate how the exclusion of aerosols affects retrieved NO₂ VCDs. Changes in NO₂ VCD are determined by CRF, VCD_{cr} and VCD_{cl} (Eq. 2). Excluding aerosols leads to a general increase in CRF (and thus the weight of VCD_{cl}) that is compensated for by a decrease in VCD_{cl} (due to an increase in CP). The change in VCD_{cr} is more complex in sign, and to a lesser extent in magnitude. In most situations, the

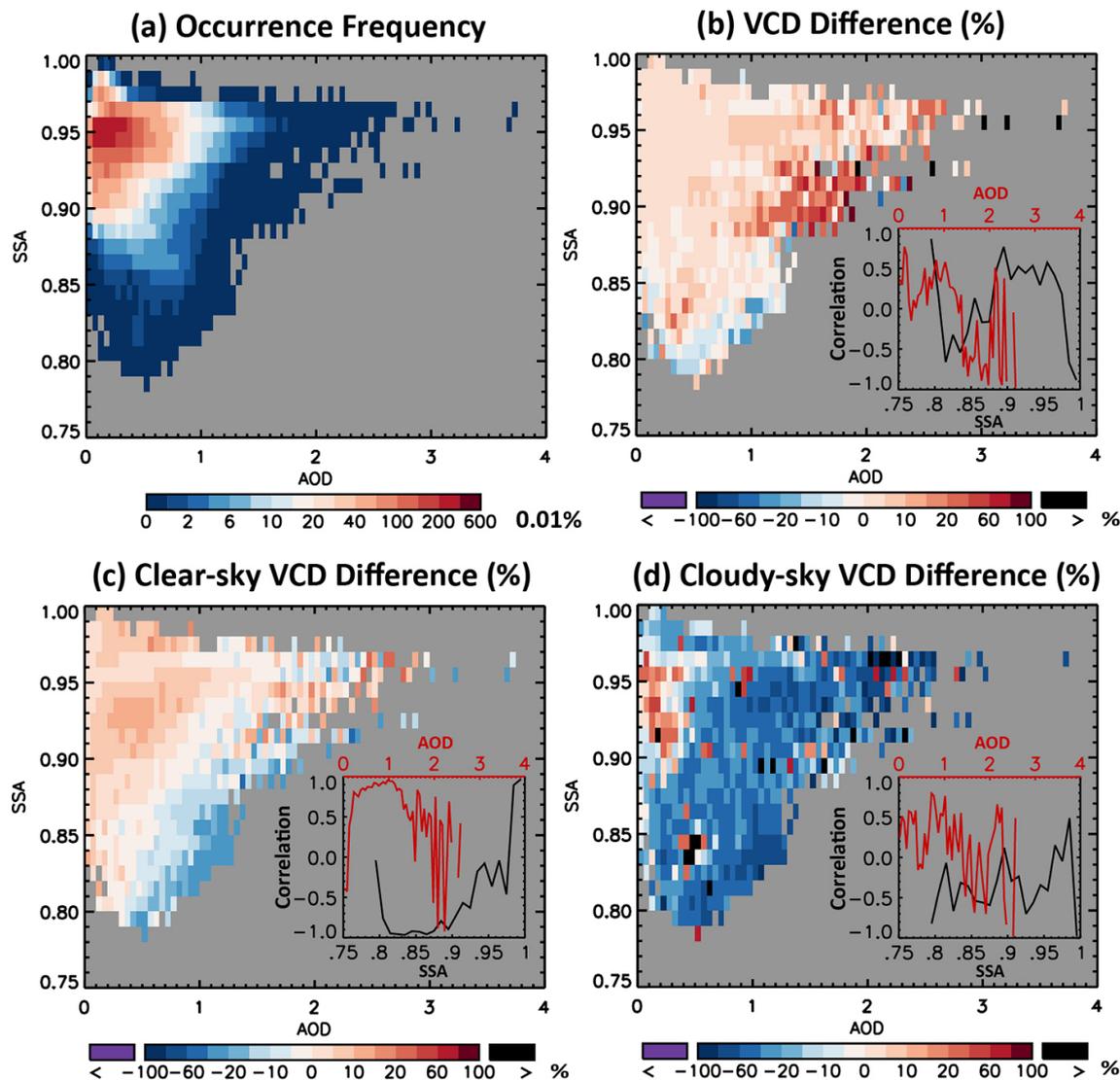


Figure 10. (a) Frequency of occurrence from all months and grid cells for each bin of AOD (x axis, bin size 0.05) and SSA (y axis, bin size 0.01). Data are sampled from valid pixels of case REF. (b) Percentage changes in NO₂ VCDs from case REF to SRF averaged over all data in each bin of AOD and SSA; also embedded are the correlations between percentage VCD changes and AOD values as a function of SSA values (black line) and between VCD changes and SSA values as a function of AOD values (red line). Similar panels are drawn for (c) clear-sky VCDs and (d) cloudy-sky VCDs.

magnitude of VCD_{cr} is much smaller than VCD_{cl} (Fig. S5 in the Supplement); this is because clouds are normally above the NO₂-concentrated layer, producing a “screening” effect on the radiation absorbed by NO₂. These factors explain the distinctive patterns of changes in NO₂ VCD_{cr}, VCD_{cl} and VCD from case REF to AER. For example, for AOD \sim 1.5 and SSA \sim 0.90, both VCD_{cr} and VCD_{cl} are reduced from case REF to AER, while the VCD is enhanced because CRF is much enhanced and VCD_{cl} greatly exceeds VCD_{cr}.

In summary, inclusion or exclusion of aerosols has distinctive effects on three independent factors, CRF (or CF), VCD_{cr}, and VCD_{cl} (or CP). Section 3.5 will show the ef-

fect of excluding aerosols on the choice of “valid” data based on the CRF criterion. It follows that an implicit treatment of aerosols cannot fully account for the complex influences of aerosols on retrieved NO₂ VCDs.

3.4 Coupled effects of aerosols and surface reflectance on NO₂ retrieval

This section evaluates the coupled effects of perturbing aerosols and surface reflectance on retrieved NO₂ VCDs. For this purpose, we use case S_A, which simultaneously adopts the OMI albedo from DOMINO v2 (following case SRF) and excludes aerosol information (following case AER) in

the retrieval process. For this case, cloud parameters are also re-derived.

Figure 5 (fourth row) shows that, at an annual scale, case S_A leads to higher NO₂ VCDs than case REF over central eastern China, southwestern China, Tibet and Qinghai. NO₂ VCDs are reduced by about 10 % over parts of the NCP. Seasonal dependence is large for the NO₂ changes. Reductions over the NCP are most significant in spring (by 10–40 %), followed by winter and fall. In summer, case S_A leads to general NO₂ enhancements by 0–40 % over most of China, including the NCP. Over the Sichuan Basin, NO₂ VCDs in fall are increased by 20–40 % from case REF to S_A, while the signs of change are more location-dependent in winter.

Figure 5 (fourth and fifth rows) shows that although case S_A follows the DOMINO v2 assumptions regarding surface reflectance and aerosols, in general it does not explain the differences between case REF and DOM. Therefore, other factors are also important in differentiating case REF from DOM. They include NO₂ vertical profiles, cloud parameters, the pixel-specific radiative transfer calculation, and air pressure (Lin et al., 2014b).

Figure 11 further shows that the effect of changing surface reflectance on retrieved NO₂ VCDs interacts with the effect of excluding aerosol information. The figure presents the difference between (SRF – REF) + (AER – REF) and (S_A – REF) as a percentage fraction of REF. Positive values occur over most of China; i.e., the effect of simultaneously changing surface reflectance and aerosols is smaller than the summed effect of changing the two parameters individually. The magnitude of differences greatly depends on seasons and locations, and is largest in winter (by 5–20 % over much of eastern China) and smallest in summer. These differences reflect the nonlinear influences of individual factors in retrieving NO₂, such that the effect of a particular parameter depends on other parameters (Lin et al., 2014b).

3.5 Discussion on the sampling of “valid” pixels and implications for analysis of pollution severity

Figure 12a shows the number of days per month with “valid” pixels based on the current criteria for cloud cover, snow/ice cover, and row-anomaly contamination. Averaged over China, fewer than half of the calendar days are included for analysis. For most of northern China, about 12–20 days per month of 2012 are selected. The number of available days per month is below 8 over much of the south due to higher cloud coverage.

Changes in aerosols and surface reflectance both have a consequence for CF and CRF. The changes in CF and CRF in turn affect the number of retrieved NO₂ results that are determined “valid” under the usual criterion of CRF < 50 %. With an implicit aerosol treatment (AER, S_A, and DOM), the CRF < 50 % criterion designates pixels as invalid when more than half of the TOA radiance is from the combination of clouds and aerosols. With an explicit aerosol treat-

ment (REF and SRF), however, the CRF < 50 % criterion means that pixels are excluded if the clouds’ contribution to the TOA radiance exceeds 50 %. Thus, the soundness of this criterion depends on how well the aerosol optical effects are quantified in the retrieval process, a critical factor in our explicit and physically more realistic aerosol treatment.

Case SRF has similar data coverage as case REF (Fig. 12b). Compared to case REF, case AER discards an additional 2–10 days per month (or 15–60 %) over most of eastern China because the aerosol-affected CRF exceeds 50 % (Fig. 12c). This is consistent with the loss of 25 % found by Lin et al. (2014b). Cases S_A and DOM also have much smaller numbers of valid days than does case REF (Fig. 12d and e).

For analyses of all retrievals in Sects. 3.1–3.4, we have evaluated retrieval results from the same set of pixels determined by case REF. In practice, however, information is not available on the pixels that could have been included in certain retrieval approaches but are instead discarded in other approaches. In such situations, a question is raised whether selecting different sets of satellite pixels will affect the evaluation of pollution severity and its subsequent applications. Of particular interest here is the loss of valid data by not explicitly accounting for aerosol optical effects (case AER vs. REF).

Figure 9 contrasts NO₂ VCDs in case AER derived from pixels “valid” in case REF (solid red lines) vs. from pixels “valid” in case AER (dashed red lines). Average VCDs for various representative regions of China are shown. A detailed horizontal distribution of their differences can be inferred by contrasting Fig. 5 (third row) and Fig. S6 in the Supplement (third row). Figure 9 shows a significant reduction in retrieved NO₂ VCDs by using the reduced set of “valid” data from case AER relative to the set determined by case REF. The reduction is about 10 % averaged over eastern China, but exceeds 50 % in some regions and months. Similar results are found for case DOM that also treats aerosol optical effects implicitly (Fig. 9, solid green vs. dashed green lines). These results arise because high aerosol loadings are often coincident with high NO₂ pollution, due to similar emission sources and meteorological influences. Discarding pixels with high aerosols (that lead to significant increases in CRF from case REF to AER and DOM) tends to exclude high NO₂ pollution situations as well (Lin et al., 2014b). This sampling bias is critical for determining the severity of NO₂ pollution from satellite remote sensing.

Explicit inclusion of aerosols in the retrieval process (as in case REF or POMINO) will avoid the sampling bias caused by removal of strong-pollution situations, especially if aerosol information can be adequately constrained. In publishing the POMINO data, we elect to include pixels with high aerosol loadings, although we note that these pixels may contain larger uncertainties for NO₂ than those with low aerosol content, since we do not have fully accurate aerosol information. This choice is supported in part by our previ-

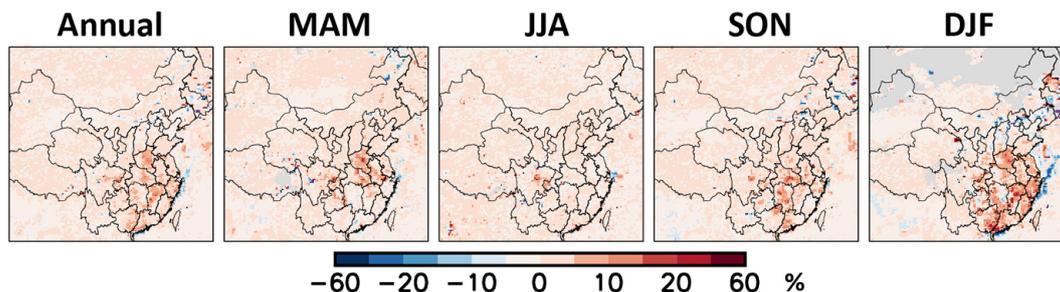


Figure 11. Differences in NO₂ VCDs between (SRF – REF) + (AER – REF) and (S_A – REF) as a percentage fraction of REF. Provincial boundaries of China are shown. Data are sampled from valid pixels of case REF. Values outside the upper (lower) bound of color intervals are shown in black (purple). Missing values are shown in grey. Color intervals are nonlinear to better present the data range; an interval without labeling represents the mean of the adjacent two intervals.

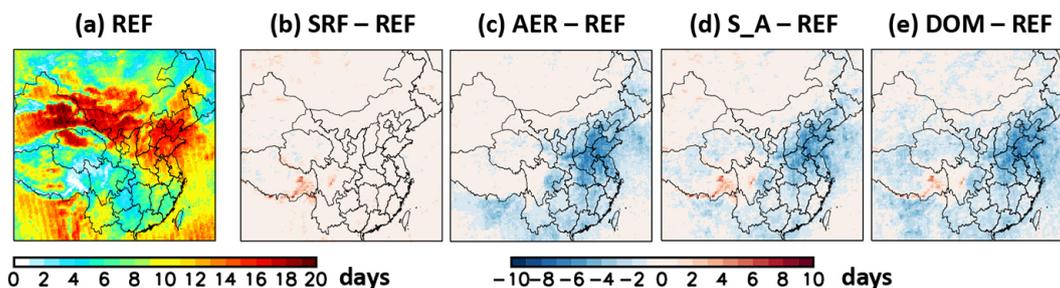


Figure 12. (a) Number of days per month of 2012 with valid pixels in case REF on a 0.25° long. × 0.25° lat. grid. (b) Changes in the number of days per month with valid pixels from REF to other cases. Provincial boundaries of China are shown.

ous comparisons against MAX-DOAS NO₂ data (Lin et al., 2014b) and by the fact that the AOD values are relatively well constrained by MODIS on a monthly basis. We publish the AOD and SSA values together with NO₂ VCDs, so that users can choose whether or not to include high-aerosol situations.

4 Impacts of aerosols and surface reflectance on OMI-based NO_x emission constraint

Sections 3.2–3.5 show the sensitivity of retrieved NO₂ VCDs to various treatments of aerosols and surface reflectance. This section further analyzes the influences on subsequent NO_x emission constraint, a popular application of OMI data. NO₂ VCDs retrieved via different approaches are used to optimize NO_x emissions. We focus on anthropogenic emissions in China. As discussed in Sect. 3.5, the set of “valid” OMI pixels depends on the retrieval approach. Thus, this section also identifies the effects of selecting various sets of “valid” pixels.

Top-down emissions are derived for individual months in 2012 on a 0.25° long. × 0.25° lat. grid, by scaling the a priori emissions (assumed in GEOS-Chem) with the ratio of OMI-derived to CTM-modeled NO₂ VCDs (Martin et al., 2003). Model NO₂ values (at 0.667° long. × 0.5° lat.) are sampled at times and locations coincident with valid OMI pixels, and are then re-mapped to the 0.25° long. × 0.25° lat. grid; see

Fig. S7 in the Supplement for the horizontal distribution of model NO₂. The scaling approach here assumes local mass balance (Leue et al., 2001). It does not fully account for the horizontal transport of NO_x or the dependence of NO_x lifetime on emissions (Turner et al., 2012); these two factors are likely to result in a small difference (within 10 %) in top-down total Chinese emissions (Lin, 2012). Also, the same scaling is applied to both anthropogenic and natural emissions. Lin (2012) presents an approach that takes advantage of the distinctive seasonality in various emission sources to better distinguish anthropogenic from natural sources. They found similar changes (from a priori to top-down) in anthropogenic and natural emissions when summed over China.

The a posteriori emissions are calculated as a weighted average of a priori and top-down emissions, by assuming normal distributions of errors in these emissions. Following Lin (2012), errors in anthropogenic emissions are taken as 60 % for a priori and 52 % for top-down for combined errors in model simulations (~40 %, Lin et al., 2012; Yan et al., 2014), satellite NO₂ retrievals (~30 %, Boersma et al., 2011; Lin et al., 2014b), and emission inversion procedures (~12 %, Lin, 2012). The same errors are assigned to all grid cells, following Lin (2012). This leads to an error of 39 % in the a posteriori emissions. Although the actual errors may be larger for individual locations, there is no such detailed information for emission constraint.

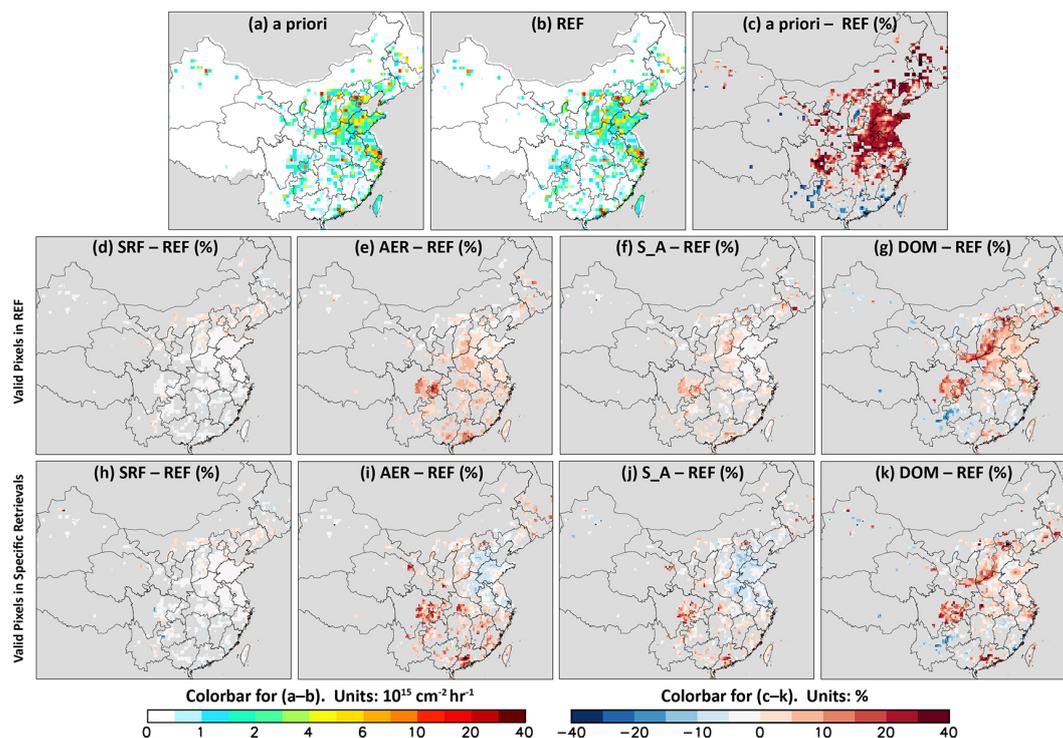


Figure 13. (a) The a priori anthropogenic emissions of NO_x on a 0.25° long. × 0.25° lat. grid. (b) The a posteriori emissions constrained by case REF. (c) Differences between a priori and REF as a percentage fraction of REF. (d–g) Changes in a posteriori emissions from REF to other cases as a percentage fraction of REF; OMI pixels are selected only when valid in case REF. (h–k) are similar to (d–g) but with respect to pixels valid in individual retrieval approaches. In (c–k), areas with emissions below $0.5 \times 10^{15} \text{ cm}^{-2} \text{ h}^{-1}$ are masked. Provincial boundaries of China are shown. Values outside the upper (lower) bound of color intervals are shown in black (purple). Color intervals are nonlinear to better present the data range; an interval without labeling represents the mean of adjacent two intervals.

Figure 13b shows the horizontal distribution of annual mean a posteriori emissions in 2012 derived from case REF. The spatial pattern is close to that of the a priori emissions (Fig. 13a) with largest values located at cities. The a posteriori emissions are lower than a priori values by 0–40 % over most of eastern China, with enhancements at many high-emission hotspots and over the southern coastal provinces (Fig. 13c).

Table 3 shows that for China as a whole, the a posteriori emissions in case REF are 9.05 TgN yr^{-1} , about 9.3 % smaller than the a priori emissions at 9.78 TgN yr^{-1} . For eastern China (101.25–126.25° E, 20–46° N), the REF a posteriori emissions for 2012 total 8.43 TgN yr^{-1} , about 19 % higher than the estimate for 2006 by Lin (2012). This increase likely reflects the recent growth of NO_x pollution in China after its recovery from the economic downturn (Lin and McElroy, 2011).

Figure 13h–k present the percentage differences in a posteriori emissions comparing results from case REF to those from the other retrieval cases. OMI pixels are selected when they are deemed to be “valid” according to criteria specific to individual retrieval cases; here each retrieval case has a distinctive set of “valid” pixels. Locations with emissions

lower than $0.5 \times 10^{15} \text{ cm}^{-2} \text{ h}^{-1}$ (totaling 11 % of Chinese emissions) are masked to highlight the polluted areas. Overall, the magnitude of emission changes is highly region-dependent. Across China with few exceptions, case SRF produces emissions similar to REF. Differences are much larger between cases REF and AER. Compared to case REF, case AER produces lower emissions by 5–10 % over the NCP but higher emissions by 5–30 % over many other regions; the greatest enhancements occur over the Sichuan Basin and the PRD. Compared to REF and AER, case S_A results in smaller emissions over the NCP. Case S_A produces emissions slightly larger than case REF in the Sichuan Basin, PRD, and many other southern provinces. By comparison, case DOM (exactly the same as DOMINO v2) leads to larger a posteriori emissions than does case REF, by 0–40 % over most of China. Table 3 shows that total Chinese emissions in all cases are within 3 % of values from case REF; this closeness is due mainly to compensation between region- and time-dependent positive and negative differences during spatiotemporal averaging.

Figure 13d–g show the percentage differences in a posteriori emissions comparing case REF to other retrieval cases, but now based on a single set of “valid” pixels de-

Table 3. A priori, top-down and a posteriori emissions of NO_x*

	China (80–130° E, 20–53° N)					Eastern China (101.25–126.25° E, 20–46° N)				
	REF	SRF	AER	S_A	DOM	REF	SRF	AER	S_A	DOM
A priori	9.78	9.78	9.78	9.78	9.78	9.11	9.11	9.11	9.11	9.11
With OMI data sampled at valid pixels specific to each case										
Top-down	8.50	8.20	8.63	8.21	8.75	7.92	7.65	8.03	7.65	8.19
A posteriori	9.05	8.88	9.12	8.88	9.19	8.43	8.27	8.49	8.27	8.58
With OMI data sampled at valid pixels by REF										
Top-down	8.50	8.23	8.91	8.53	8.80	7.92	7.68	8.33	7.97	8.27
A posteriori	9.05	8.90	9.28	9.06	9.22	8.43	8.29	8.66	8.46	8.63

* Units: TgN yr⁻¹

terminated by case REF. Locations with emissions lower than $0.5 \times 10^{15} \text{ cm}^{-2} \text{ h}^{-1}$ are masked to highlight the polluted areas. Case SRF leads to similar emissions to case REF. Case AER exceeds REF by 0–30 % over most regions, but with reductions by 0–5 % over parts of northern China. Case S_A is also larger than REF over much of China. Case DOM exceeds REF by 0–40 % especially over the Sichuan Basin and parts of northern China.

Differences between Fig. 13e–g and i–k are apparent – a result of different pixel sampling. The differences over the NCP indicate that, since the implicit treatment of aerosols in cases AER, S_A and DOM leads to exclusion of situations with high aerosol loadings (and coincidentally high NO₂ pollution; see Sect. 3.5), there is a consequent underestimate in the a posteriori emissions.

Figure 14a, d and g shows the contrasts between cases REF and AER for the ratio of maximum to minimum monthly a posteriori emissions. Locations are shown only with emissions greater than $0.5 \times 10^{15} \text{ cm}^{-2} \text{ h}^{-1}$. For case REF (Fig. 14a), the max/min ratios range from 1 to 10, but are about 2–5 at most locations. Figure 14d (based on REF-determined valid pixels) shows enhancements in the max/min ratio from case REF to case AER over most of China. By comparison, Fig. 14g (based on case-specific sets of valid pixels) shows that case AER leads to much larger max/min ratios over most of southern China, while the changes in max/min ratios are more location-dependent over the NCP. Figure 14 further contrasts cases REF and AER for the months of maximum and minimum emissions. For case REF, minimum emissions often occur in May–September in the NCP, in contrast to December–March in many southern areas (Fig. 14b). Maximum emissions occur in November–February over most of China (Fig. 14c). Case AER leads to location-dependent shifts in the months of maximum and minimum emissions (Fig. 14e, f, h, and i); furthermore, these shifts are more apparent when OMI pixels are selected with respect to individual retrieval approaches (Fig. 14h and i). These differences highlight the seasonal and spatial depen-

dence of constrained emissions on the NO₂ retrieval approaches.

5 Conclusions

This paper presents an improved retrieval, the POMINO algorithm, of tropospheric NO₂ VCDs over China in 2012, by explicitly accounting for temporally and spatially varying aerosol optical effects and surface reflectance anisotropy. Prerequisite cloud optical parameters (CF and CP) are retrieved with the same treatments for aerosols and surface, thus eliminating any algorithm inconsistency in retrieving NO₂ and clouds. Our NO₂ retrieval is focused on calculations of tropospheric AMFs, taking the tropospheric SCDs from DOMINO v2. Our cloud retrieval starts from the SCDs of O₂–O₂ in OMCLDO2 v3. Aerosol vertical profiles and optical properties are first taken from GEOS-Chem simulations, with subsequent AOD constraints based on monthly MODIS Aqua data. Surface reflectance data over land are taken from the MODIS BRDF product. Retrievals are performed for individual OMI pixels via online radiative transfer calculations without the need for look-up tables, thanks to the newly parallelized AMFv6 code driven by LIDORT v3.6. Sensitivity tests are performed to evaluate the effect of an explicit (as opposed to an implicit) treatment of aerosols and the effect of changes in surface reflectance characteristics on retrieved NO₂ VCDs. Additional analyses are done for subsequent OMI-based NO_x emission constraints. Results are presented as monthly mean values on a $0.25^\circ \text{ long.} \times 0.25^\circ \text{ lat.}$ grid.

POMINO NO₂ VCDs undergo strong seasonal and spatial variability. Large NO₂ VCDs are located in eastern China due to significant anthropogenic emissions. On an annual basis, NO₂ VCDs vary from $15\text{--}25 \times 10^{15} \text{ cm}^{-2}$ over the NCP to below 10^{15} cm^{-2} over much of western China. Over the polluted regions, NO₂ VCDs reach maxima in winter and minima in summer, due mostly to the seasonal variation in

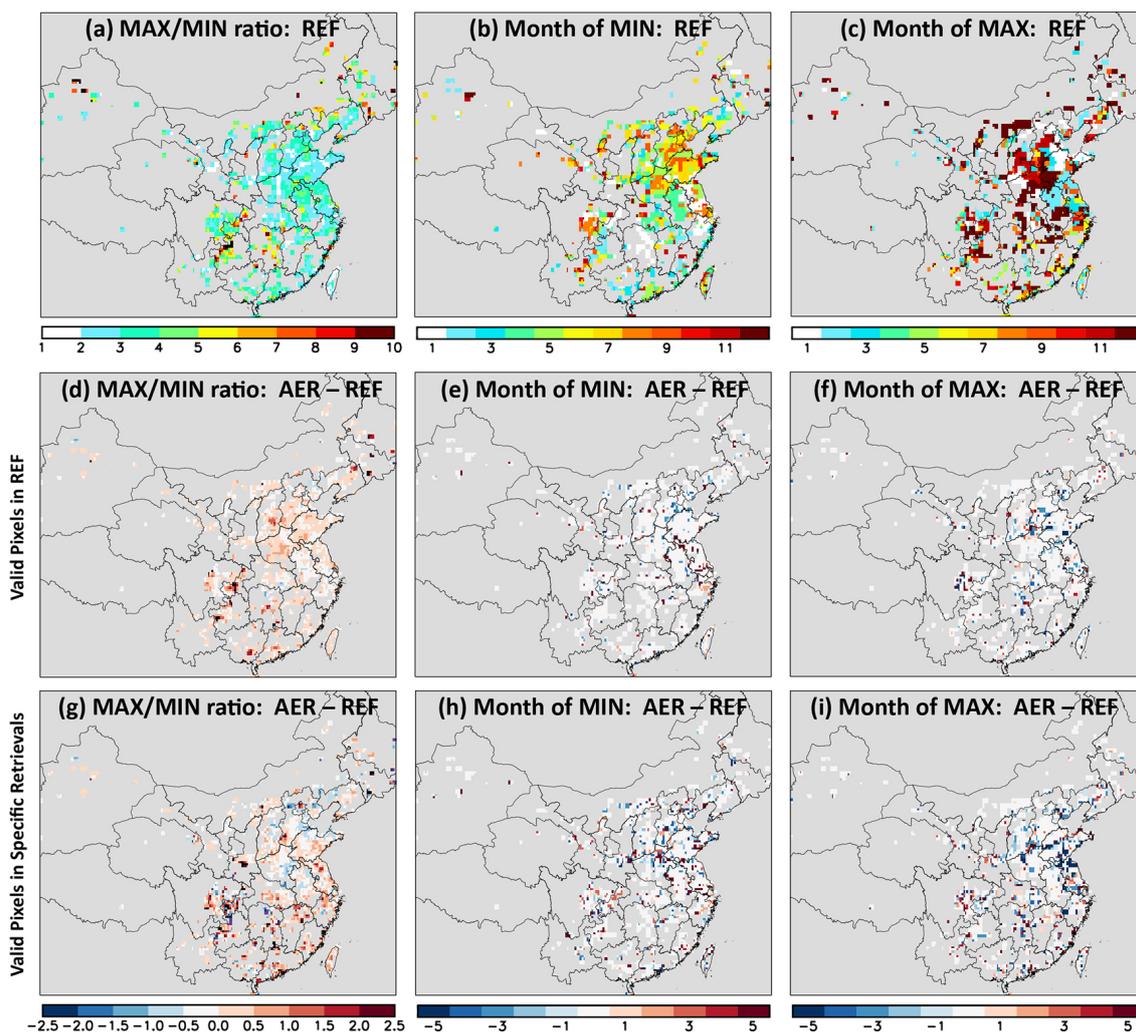


Figure 14. (a) The maximum to minimum ratio for monthly a posteriori anthropogenic emissions in case REF, and the months of (b) minimum and (c) maximum emissions. (d–f) Differences between cases AER and REF; OMI pixels are selected only when valid in case REF. (g–i) are similar to (d–f) but with respect to pixels valid in individual retrieval cases. Provincial boundaries of China are shown. Values outside the upper (lower) bound of color intervals are shown in black (purple). In (d–i), areas with emissions below $0.5 \times 10^{15} \text{ cm}^{-2} \text{ h}^{-1}$ are masked. In (e), (f), (h) and (i), the month proceeds in loop; i.e., November (January) is 1 month behind (ahead of) December.

NO_x lifetime. Over cleaner regions (e.g., much of western China), NO₂ VCDs peak in summer with minima in winter, because strong natural emissions in summer overcompensate for the short NO₂ lifetime. The maximum to minimum ratio in monthly mean NO₂ VCDs is about 3.6 for regions east of 101.25° E as a whole and about 1.4 over the west. A POMINO-based NO_x emission constraint leads to a posteriori Chinese anthropogenic emissions at 9.05 TgN yr^{-1} , an increase from 2006 (Lin, 2012) by about 19%.

In one sensitivity test, we re-retrieved clouds and NO₂ VCDs, by adopting the monthly climatological OMI albedo data (OMLER v1) in place of the MODIS BRDF data. Surface reflectance is greatly enhanced over most of eastern China and reduced over the west. Changes in surface reflectance result in changes of opposite sign in CF with more

complex effects on CP. As a consequence, annual mean NO₂ VCDs are decreased by 5–15% over central and southern China, with enhancements over many other regions. The reductions in southern China are largest (15–40%) in fall, but are replaced by slight enhancements in summer. On an annual basis, changes in surface reflectance have a small effect (within 5%) on constrained NO_x emissions over most Chinese locations with anthropogenic emissions greater than $0.5 \times 10^{15} \text{ cm}^{-2} \text{ h}^{-1}$.

In another sensitivity test, an implicit treatment of aerosols mimics the traditional algorithms by excluding aerosol information in retrieving clouds and NO₂ VCDs. The implicit treatment greatly enhances effective CF (the spatial correlation is 0.75 between neglected annual mean AOD values and CF enhancements), CRF (correlation 0.82), and CP (corre-

lation 0.19). The low correlation for CP highlights the complexity of aerosol effects. Changes in NO₂ VCDs from an explicit to an implicit treatment of aerosols depend in a complex manner on AOD, SSA, heights of aerosols relative to NO₂, and other factors; these factors affect CRF, VCDcl and VCDcr that determine NO₂ VCDs. The annual mean NO₂ VCDs are enhanced by 15–40 % over much of eastern China. The seasonal and spatial variability of NO₂ changes is apparent. NO₂ VCDs are reduced by 10–20 % over parts of the NCP in spring and over northern China in winter, whereas these reductions are replaced by general enhancements in summer and fall. The effect on subsequently constrained annual NO_x emissions varies between –5 and 30 % over eastern China with apparent seasonal and regional dependence.

For the usual criterion of CRF < 50 % to select “valid” pixels, the large enhancements in CRF due to an implicit aerosol treatment result in significant reductions in the number of days with “valid” pixels over polluted regions (by 2–10 days per month, or 15–60 %, on average). This also leads to a likely sampling bias due to the exclusion of high-aerosol days that often experience high NO₂ pollution, consistent with the findings of Lin et al. (2014b).

The effect of an implicit aerosol treatment on retrieved NO₂ VCDs also interacts with the effect of changes in surface reflectance. NO₂ changes obtained by simultaneously excluding aerosol information and adopting the OMI albedo are smaller than the overall NO₂ changes obtained by making these two adjustments individually and summing the results. Although these adjustments (no aerosols + OMI albedo) are present in the DOMINO v2 approach, the resulting NO₂ VCDs still differ from those of DOMINO v2, indicating the importance of differences between other factors assumed in the retrieval processes (e.g., NO₂ profile shape, atmospheric profile in O₂–O₂ retrieval, pixel-by-pixel radiative transfer calculation vs. look-up table, and air pressure).

There are no sufficiently comprehensive independent measurements available to systematically evaluate the various NO₂ retrieval approaches. Current MAX-DOAS measurements are very limited over China, with few sites and short operation periods (Irie et al., 2012; Ma et al., 2013; Hendrick et al., 2014; Kanaya et al., 2014). In situ measurements are rare for vertical profiles of aerosols and NO₂. Our results show that the effects of aerosols and surface reflectance are highly season- and location-dependent. This clearly indicates the need for a comprehensive measurement network to validate satellite data. Nonetheless, our present study and that of Lin et al. (2014b) point the way forward for a physically more realistic NO₂ retrieval by explicit inclusion of aerosol effects.

Currently, our POMINO NO₂ data are available for 2004–2013 (<http://www.atmos.pku.edu.cn/acm/acmModel.html#POMINO>). Daily level-3 data on a 0.25° long. × 0.25° lat. grid are provided on the webpage for general users, and level-2 data (including averaging kernels) can be provided for advanced users upon request. We simultaneously provide daily AOD, SSA and surface reflectance (MODIS BRF over land and turbid coastal ocean and OMLER v3 albedo over the open ocean) data for users. We elect to include daily NO₂ data with high aerosol loadings to reduce the potential sampling bias, in line with our choice of an explicit aerosol treatment and as supported by Lin et al. (2014b). We note that these NO₂ data with high aerosol pollution may be subject to larger uncertainties, since aerosols are not fully accurately constrained by observations. Users can make their own judgment on whether to include the NO₂ data with high aerosol pollution. With more aerosol data from satellite, ground and in situ measurements in the future, aerosol optical effects can be better constrained to reduce the associated uncertainties in NO₂ retrievals.

Our parallelized LIDORT-driven AMFv6 code is also available for public use. With 16 computational cores (Intel (R) Xeon (R) CPU X7550 at 2.00 GH) running in parallel, it takes about 3 h of wall-clock time to retrieve both cloud parameters and NO₂ VCDs over China for a month; the retrieval process includes pixel-specific radiative transfer calculations and explicit treatments of aerosols and surface reflectance anisotropy. The excellent scalability of our parallelized code means that additional computational cores can be employed to further speed up the retrieval process. Such retrieval efficiency enables a fast global retrieval that will be particularly important for future fine-resolution satellite instruments such as TropOMI (which is expected to have a data rate ~ 8 times that of OMI) and GEMS (which will be onboard a geostationary satellite with hourly measurements at a horizontal resolution of 5 × 15 km²).

Appendix A: GEOS-Chem simulations

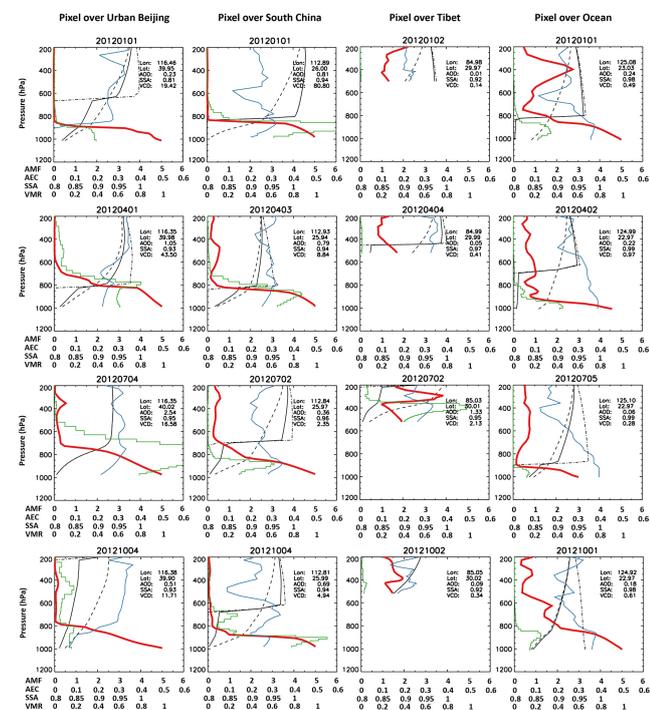


Figure A1. Vertical profiles of NO₂ volume mixing ratios (VMRs, thick red lines), aerosol extinction coefficients (AECs, green lines), SSA (blue lines), AMFs (black lines), clear-sky AMFs (black dashed lines), and cloudy-sky AMFs (black dotted dashed lines) at the beginning of January, April, July and October at representative pixels in four areas: urban Beijing, southern China, Tibet, and the oceans. The VMRs are normalized to the maximum VMR of the vertical profile. Embedded in each panel are longitude, latitude, AOD, SSA and retrieved NO₂ VCDs (in 10^{15} cm^{-2}). Here, “20120101” in the title of the top-left panel means 1 January 2012, and so on.

We adopted vertical profiles of NO₂ and aerosol optical parameters from GEOS-Chem simulations on a $0.667^\circ \text{ long.} \times 0.5^\circ \text{ lat.}$ grid. We used the latest version, 9-02, which is an update of version v8-03-02 employed by Lin et al. (2014b). Here we summarize the model updates relevant to our retrieval study. The reader is referred to Lin et al. (2014b) for a general model description.

GEOS-Chem simulates spatiotemporally varying concentrations of NO_x and other gaseous and aerosol species. It also outputs optical properties of various aerosol types, including secondary inorganic aerosols (sulfates, nitrates, and ammoniums), organic aerosols, black carbon, dust and sea salts. Vertical profiles and type-specific optical properties of aerosols are available for our NO₂ retrieval, including size distributions, refractive indices, AOD, SSA, phase function, and hygroscopic growth rates (Drury et al., 2010). The simulations have been conducted from 2004 to 2013 to facilitate our NO₂ retrieval.

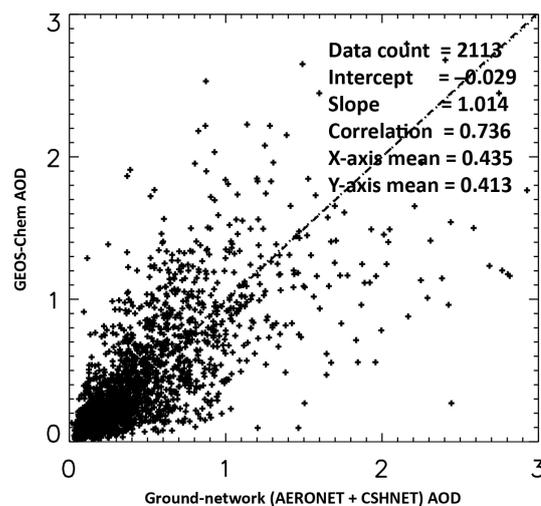


Figure A2. Scatterplot of AOD at 550 nm from GEOS-Chem simulations vs. ground networks (AERONET + CSHNET). Model AOD has been adjusted by MODIS AOD on a monthly basis (see Appendix B).

For Chinese anthropogenic emissions, we used the latest MEIC inventory for NO_x and other gaseous species with a base year of 2008 (www.meicmodel.org). The inventory is an update of the previous INTEX-B inventory. We followed Huang et al. (2012) for anthropogenic emissions of ammonia. Other emission setups are the same as those in Lin et al. (2014b). In China, anthropogenic emissions of NO_x have varied significantly in recent years. Thus, we calculated the ratios of DOMINO v2 NO₂ VCDs in 2008 to other years on a $0.25^\circ \text{ long.} \times 0.25^\circ \text{ lat.}$ grid, and then applied the ratios to scale NO_x emissions to the respective years. Such scaling is tantamount to a first-order emission adjustment to reduce the effect of emission-dependent NO_x lifetime on emission constraints in Sect. 4.

We implemented seasonality for Chinese residential emissions to reflect its current residential heating scheme (Streets et al., 2003). Residential emissions are caused by stove operation. As a modification of the piecewise method by Streets et al. (2003), we assumed that residential stoves operate 16 h per day (h d^{-1}) for monthly mean 2 m air temperature below 0°C , 3 h d^{-1} for temperature above 10°C , and $3\text{--}16 \text{ h d}^{-1}$ in between the two temperatures (linear interpolation). Monthly mean air temperature is taken from the GEOS-5 meteorological fields and averaged over 2005–2009. As such, residential emissions are largest in winter and smallest in summer. The seasonality is not important for NO_x since residential emissions contribute 5.6% of total emissions only (Zhang et al., 2009). By comparison, seasonality has a large impact on black carbon, for which pollutant residential sources contribute 55% of total anthropogenic emissions in China (Zhang et al., 2009). The emission seasonality greatly affects the SSA of total aerosols.

GEOS-Chem employs a non-local scheme to simulate vertical mixing in the boundary layer (Lin and McElroy, 2010) and a modified relaxed Arakawa–Schubert scheme for convection (Rienecker et al., 2008). GEOS-Chem captures vertical profiles of NO₂ and ozone over the United States from aircraft measurements (Lin and McElroy, 2010). The same schemes are used for NO₂ and aerosols – thus, errors in the height of aerosols relative to NO₂ may be smaller than errors in their absolute heights. This is important for reducing the sensitivity of retrieved NO₂ VCDs to errors in the vertical profiles of aerosols, given the lack of observational constraints on their profiles. Figure A1, as an illustration, presents vertical profiles of NO₂ volume mixing ratios (VMRs), aerosol extinction coefficients (AECs), SSA, AMFs, clear-sky AMFs, and cloudy-sky AMFs at the beginning of January, April, July and October at representative pixels in four regions: urban Beijing, southern China, Tibet, and the oceans. Figure A1 shows that over the polluted eastern China, NO₂ and aerosols are often collocated in the boundary layer. Over Tibet and the oceans (clear air), there exist large amounts of NO₂ above the aerosol layer. A comparison of our NO₂ profiles with those in DOMINO v2 (based on the TM4 model) is presented in Fig. 3 of Lin et al. (2014b).

Appendix B: Using MODIS AOD to constrain GEOS-Chem AOD

We used MODIS Aqua AOD data from the MYD04 Collection 5.1 level-2 dark-target product (Remer et al., 2008) to constrain model AOD on a monthly basis. Adjustments were done at a wavelength of 550 nm.

We first projected the level-2 MODIS AOD data to the GEOS-Chem 0.667° long. × 0.5° lat. grid for each day, in order to perform a subsequent calculation of monthly means. The derived monthly mean MODIS data are subject to missing values, due to stringent cloud screening and difficulties in retrieving AOD above snow-/ice-covered or bare/deserted lands. Thus, we conducted temporal and spatial interpolation upon the monthly mean data. We used the average of monthly mean values from previous and successive months to fill the missing value in a particular month (first the adjacent 2 months, then the adjacent 4 months, and so on); the missing value was filled when at least two adjacent months with valid data were found. For the temporal interpolation, we made use of monthly data from 2004 to 2013. For grid cells containing missing values after the temporal interpolation, we performed subsequent spatial interpolation in a similar manner: first the adjacent 5 × 5 = 25 grid cells were used, and then the adjacent 7 × 7 = 49 grid cells, and so on. A missing value was filled when at least 15 adjacent grid cells are deemed to be valid. Figure S8 in the Supplement presents the AOD values before and after this interpolation.

We note here that the latest MODIS Collection 6 product provides a data set merging results from the “dark-target” and “deep-blue” algorithms; this merger greatly reduces the number of missing values (Sayer et al., 2014). The dark-target algorithm has changed little from Collection 5.1 to 6. In the future, we plan to adopt this new data set to minimize the need for data interpolation.

We used the data-filled monthly mean MODIS AOD data to adjust GEOS-Chem model results. To facilitate the adjustment, a corresponding monthly mean model data set was made. We selected model AOD at times coincident with pre-data-filling MODIS data, and then went through the same temporal and spatial interpolation process to generate a corresponding data-filled monthly mean data set.

For a given month, differences between the data-filled monthly mean MODIS and model AOD data sets were used to adjust the original model results for each day of the month. We employed two adjustment methods. The primary method calculates the ratio of data-filled monthly mean MODIS AOD to data-filled model AOD and then applies this ratio to scale model results for all days of the respective month. This scaling method can lead to unrealistically large AOD for certain grid cells. In this eventuality, we simply calculated the arithmetic difference between data-filled monthly mean MODIS AOD and data-filled model AOD, and applied this difference to the original model AOD for all days of the month. This secondary adjustment was applied when the scaling factor exceeded a threshold of 3, or when the monthly mean (from all days) of the original model AOD multiplied by the scaling factor exceeded a value of 1.5.

To validate our model adjustment, we compared the post-adjustment model AOD with independent AOD data from the two ground measurement networks AERONET (Holben et al., 1998) and CSHNET (Xin et al., 2007). These two networks encompass a total of 19 sites across China, with various ground conditions and aerosol loadings; see Lin et al. (2014c) for detailed descriptions of these ground networks. The scatterplot in Fig. A2 compares the post-adjustment model data with ground networks in 2006. The post-adjustment model AOD is consistent with ground networks, with a correlation of 0.74, a small mean bias of −0.02 (5%), a reduced major axis regression slope of 1.01 and an intercept of −0.03. We therefore conclude that the model adjustment leads to a reasonable AOD constraint.

The Supplement related to this article is available online at doi:10.5194/acp-15-11217-2015-supplement.

Acknowledgements. This research is supported by the National Natural Science Foundation of China, grants 41175127 and 41422502, and by the 973 program, grant 2014CB441302. We acknowledge the free use of NO₂ and cloud products from <http://www.temis.nl>, GMTED2010 surface elevation data from USGS, GLCNMO land use type data from ISCGM, MODIS aerosol and surface reflectance data from NASA, and several AOD measurements from AERONET. Folkert Boersma acknowledges receiving funding for this research from NWO Vidi grant 864.09.001 and from the European Community's Seventh Framework Programme under grant agreement no. 607405 (QA4ECV). The calculation was supported in part by the National Institute of Supercomputing and Network/Korea Institute of Science and Technology Information with supercomputing resources including technical support (KSC-2014-C2-034).

Edited by: G. Carmichael

References

- Acarreta, J. R., De Haan, J. F., and Stammes, P.: Cloud pressure retrieval using the O-2-O-2 absorption band at 477 nm, *J. Geophys. Res.-Atmos.*, 109, D05204, doi:10.1029/2003jd003915, 2004.
- Boersma, K. F., Eskes, H. J., and Brinksma, E. J.: Error analysis for tropospheric NO₂ retrieval from space, *J. Geophys. Res.*, 109, D04311, doi:10.1029/2003JD003962, 2004.
- Boersma, K. F., Eskes, H. J., Dirksen, R. J., van der A, R. J., Veefkind, J. P., Stammes, P., Huijnen, V., Kleipool, Q. L., Sneep, M., Claas, J., Leitão, J., Richter, A., Zhou, Y., and Brunner, D.: An improved tropospheric NO₂ column retrieval algorithm for the Ozone Monitoring Instrument, *Atmos. Meas. Tech.*, 4, 1905–1928, doi:10.5194/amt-4-1905-2011, 2011.
- Bucsela, E. J., Krotkov, N. A., Celarier, E. A., Lamsal, L. N., Swartz, W. H., Bhartia, P. K., Boersma, K. F., Veefkind, J. P., Gleason, J. F., and Pickering, K. E.: A new stratospheric and tropospheric NO₂ retrieval algorithm for nadir-viewing satellite instruments: applications to OMI, *Atmos. Meas. Tech.*, 6, 2607–2626, doi:10.5194/amt-6-2607-2013, 2013.
- Castellanos, P. and Boersma, K. F.: Reductions in nitrogen oxides over Europe driven by environmental policy and economic recession, *Sci. Rep.*, 2, 265, doi:10.1038/srep00265, 2012.
- Castellanos, P., Boersma, K. F., Torres, O., and de Haan, J. F.: OMI tropospheric NO₂ air mass factors over South America: effects of biomass burning aerosols, *Atmos. Meas. Tech.*, 8, 3831–3849, doi:10.5194/amt-8-3831-2015, 2015.
- Che, H. Z., Yang, Z. F., Zhang, X. Y., Zhu, C. Z., Ma, Q. L., Zhou, H. G., and Wang, P.: Study on the aerosol optical properties and their relationship with aerosol chemical compositions over three regional background stations in China, *Atmos. Environ.*, 43, 1093–1099, doi:10.1016/j.atmosenv.2008.11.010, 2009.
- Cooper, O. R., Parrish, D. D., Stohl, A., Trainer, M., Nedelec, P., Thouret, V., Cammas, J. P., Oltmans, S. J., Johnson, B. J., Tarasick, D., Leblanc, T., McDermid, I. S., Jaffe, D., Gao, R., Stith, J., Ryerson, T., Aikin, K., Campos, T., Weinheimer, A., and Avery, M. A.: Increasing springtime ozone mixing ratios in the free troposphere over western North America, *Nature*, 463, 344–348, doi:10.1038/nature08708, 2010.
- Dirksen, R. J., Boersma, K. F., Eskes, H. J., Ionov, D. V., Bucsela, E. J., Levelt, P. F., and Kelder, H. M.: Evaluation of stratospheric NO₂ retrieved from the Ozone Monitoring Instrument: Intercomparison, diurnal cycle, and trending, *J. Geophys. Res.-Atmos.*, 116, D08305, doi:10.1029/2010jd014943, 2011.
- Drury, E., Jacob, D. J., Spurr, R. J. D., Wang, J., Shinzuka, Y., Anderson, B. E., Clarke, A. D., Dibb, J., McNaughton, C., and Weber, R.: Synthesis of satellite (MODIS), aircraft (ICARTT), and surface (IMPROVE, EPA-AQS, AERONET) aerosol observations over eastern North America to improve MODIS aerosol retrievals and constrain surface aerosol concentrations and sources, *J. Geophys. Res.-Atmos.*, 115, D14204, doi:10.1029/2009jd012629, 2010.
- Ford, B. and Heald, C. L.: An A-train and model perspective on the vertical distribution of aerosols and CO in the Northern Hemisphere, *J. Geophys. Res.-Atmos.*, 117, D06211, doi:10.1029/2011jd016977, 2012.
- Hendrick, F., Müller, J.-F., Clémer, K., Wang, P., De Mazière, M., Fayt, C., Gielen, C., Hermans, C., Ma, J. Z., Pinardi, G., Stavrou, T., Vlemmix, T., and Van Roozendaal, M.: Four years of ground-based MAX-DOAS observations of HONO and NO₂ in the Beijing area, *Atmos. Chem. Phys.*, 14, 765–781, doi:10.5194/acp-14-765-2014, 2014.
- Holben, B. N., Eck, T. F., Slutsker, I., Tanre, D., Buis, J. P., Setzer, A., Vermote, E., Reagan, J. A., Kaufman, Y. J., Nakajima, T., Lavenu, F., Jankowiak, I., and Smirnov, A.: AERONET – A federated instrument network and data archive for aerosol characterization, *Remote Sens. Environ.*, 66, 1–16, 1998.
- Huang, X., Song, Y., Li, M., Li, J., Huo, Q., Cai, X., Zhu, T., Hu, M., and Zhang, H.: A high-ammonia emission inventory in China, *Global Biogeochem. Cy.*, 26, GB1030, doi:10.1029/2011GB004161, 2012.
- Hyer, E. J., Reid, J. S., and Zhang, J.: An over-land aerosol optical depth data set for data assimilation by filtering, correction, and aggregation of MODIS Collection 5 optical depth retrievals, *Atmos. Meas. Tech.*, 4, 379–408, doi:10.5194/amt-4-379-2011, 2011.
- Irie, H., Boersma, K. F., Kanaya, Y., Takashima, H., Pan, X., and Wang, Z. F.: Quantitative bias estimates for tropospheric NO₂ columns retrieved from SCIAMACHY, OMI, and GOME-2 using a common standard for East Asia, *Atmos. Meas. Tech.*, 5, 2403–2411, doi:10.5194/amt-5-2403-2012, 2012.
- Jiang, Z., Worden, J. R., Jones, D. B. A., Lin, J.-T., Verstraeten, W. W., and Henze, D. K.: Constraints on Asian ozone using Aura TES, OMI and Terra MOPITT, *Atmos. Chem. Phys.*, 15, 99–112, doi:10.5194/acp-15-99-2015, 2015.
- Kanaya, Y., Irie, H., Takashima, H., Iwabuchi, H., Akimoto, H., Sudo, K., Gu, M., Chong, J., Kim, Y. J., Lee, H., Li, A., Si, F., Xu, J., Xie, P.-H., Liu, W.-Q., Dzhola, A., Postlyakov, O., Ivanov, V., Grechko, E., Terpigova, S., and Panchenko, M.: Long-term MAX-DOAS network observations of NO₂ in Russia and Asia (MADRAS) during the period 2007–2012: instrumentation, elucidation of climatology, and comparisons with OMI satellite observations and global model simulations, *Atmos. Chem. Phys.*, 14, 7909–7927, doi:10.5194/acp-14-7909-2014, 2014.

- Kleipool, Q. L., Dobber, M. R., de Haan, J. F., and Lev-
elt, P. F.: Earth surface reflectance climatology from 3
years of OMI data, *J. Geophys. Res.-Atmos.*, 113, D18308,
doi:10.1029/2008jd010290, 2008.
- Lamsal, L. N., Martin, R. V., Padmanabhan, A., van Donke-
laar, A., Zhang, Q., Sioris, C. E., Chance, K., Kurosu, T.
P., and Newchurch, M. J.: Application of satellite obser-
vations for timely updates to global anthropogenic NO(x)
emission inventories, *Geophys. Res. Lett.*, 38, L05810,
doi:10.1029/2010gl046476, 2011.
- Lee, K. H., Li, Z., Wong, M. S., Xin, J., Wang, Y., Hao, W. M., and
Zhao, F.: Aerosol single scattering albedo estimated across China
from a combination of ground and satellite measurements, *J.
Geophys. Res.*, 112, D22S15, doi:10.1029/2007JD009077, 2007.
- Leue, C., Wenig, M., Wagner, T., Klimm, O., Platt, U., and Jähne,
B.: Quantitative analysis of NO_x emissions from Global Ozone
Monitoring Experiment satellite image sequences, *J. Geophys.
Res.*, 106, 5493–5505, doi:10.1029/2000JD900572, 2001.
- Levelt, P. F., van den Oord, G. H., Dobber, M. R., Malkki, A., Visser,
H., de Vries, J., Stammes, P., Lundell, J. O., and Saari, H.: The
ozone monitoring instrument, *IEEE Trans. Geosci. Remote.*, 44,
1093–1101, 2006.
- Lin, J.-T.: Satellite constraint for emissions of nitrogen oxides from
anthropogenic, lightning and soil sources over East China on
a high-resolution grid, *Atmos. Chem. Phys.*, 12, 2881–2898,
doi:10.5194/acp-12-2881-2012, 2012.
- Lin, J. T. and McElroy, M. B.: Impacts of boundary layer mixing
on pollutant vertical profiles in the lower troposphere: Impli-
cations to satellite remote sensing, *Atmos. Environ.*, 44,
1726–1739, doi:10.1016/j.atmosenv.2010.02.009, 2010.
- Lin, J.-T. and McElroy, M. B.: Detection from space of a reduction
in anthropogenic emissions of nitrogen oxides during the Chi-
nese economic downturn, *Atmos. Chem. Phys.*, 11, 8171–8188,
doi:10.5194/acp-11-8171-2011, 2011.
- Lin, J.-T., Liang, X.-Z., and Wuebbles, D. J.: Effects of Intercon-
tinent Transport on Surface Ozone over the United States:
Present and Future Assessment with a Global Model, *Geophys.
Res. Lett.*, 35, L02805, doi:10.1029/2007GL031415, 2008.
- Lin, J.-T., Nielsen, C. P., Zhao, Y., Lei, Y., Liu, Y., and McEl-
roy, M. B.: Recent Changes in Particulate Air Pollution over
China Observed from Space and the Ground: Effectiveness
of Emission Control, *Environ. Sci. Technol.*, 44, 7771–7776,
doi:10.1021/es101094t, 2010a.
- Lin, J.-T., McElroy, M. B., and Boersma, K. F.: Constraint of an-
thropogenic NO_x emissions in China from different sectors:
a new methodology using multiple satellite retrievals, *Atmos.
Chem. Phys.*, 10, 63–78, doi:10.5194/acp-10-63-2010, 2010b.
- Lin, J.-T., Pan, D., Davis, S. J., Zhang, Q., He, K., Wang, C., Streets,
D. G., Wuebbles, D. J., and Guan, D.: China's international trade
and air pollution in the United States, *P. Natl. Acad. Sci. USA*,
111, 15036–15041, doi:10.1073/pnas.1312860111, 2014a.
- Lin, J.-T., Martin, R. V., Boersma, K. F., Sneep, M., Stammes, P.,
Spurr, R., Wang, P., Van Roozendaal, M., Clémer, K., and Irie,
H.: Retrieving tropospheric nitrogen dioxide from the Ozone
Monitoring Instrument: effects of aerosols, surface reflectance
anisotropy, and vertical profile of nitrogen dioxide, *Atmos.
Chem. Phys.*, 14, 1441–1461, doi:10.5194/acp-14-1441-2014,
2014b.
- Lin, J., van Donkelaar, A., Xin, J., Che, H., and Wang, Y.: Clear-sky
aerosol optical depth over East China estimated from visibility
measurements and chemical transport modeling, *Atmos. Envi-
ron.*, 95, 257–267, doi:10.1016/j.atmosenv.2014.06.044, 2014c.
- Lin, J.-T., Liu, Z., Zhang, Q., Liu, H., Mao, J., and Zhuang,
G.: Modeling uncertainties for tropospheric nitrogen dioxide
columns affecting satellite-based inverse modeling of nitro-
gen oxides emissions, *Atmos. Chem. Phys.*, 12, 12255–12275,
doi:10.5194/acp-12-12255-2012, 2012.
- Liu, J., Kuang, W., Zhang, Z., Xu, X., Qin, Y., Ning, J., Zhou, W.,
Zhang, S., Li, R., and Yan, C.: Spatiotemporal characteristics,
patterns, and causes of land-use changes in China since the late
1980s, *J. Geogr. Sci.*, 24, 195–210, 2014.
- Lucht, W., Schaaf, C. B., and Strahler, A. H.: An algorithm
for the retrieval of albedo from space using semiempirical
BRDF models, *IEEE Trans. Geosci. Remote.*, 38, 977–998,
doi:10.1109/36.841980, 2000.
- Ma, J. Z., Beirle, S., Jin, J. L., Shaiganfar, R., Yan, P., and Wagner,
T.: Tropospheric NO₂ vertical column densities over Beijing: re-
sults of the first three years of ground-based MAX-DOAS mea-
surements (2008–2011) and satellite validation, *Atmos. Chem.
Phys.*, 13, 1547–1567, doi:10.5194/acp-13-1547-2013, 2013.
- Maasakkers, J. D.: Vital improvements to the retrieval of tropo-
spheric NO₂ columns from the Ozone Monitoring Instrument,
Eindhoven University of Technology, Eindhoven, 2013.
- Martin, R. V., Jacob, D. J., Chance, K., Kurosu, T. P., Palmer, P.
I., and Evans, M. J.: Global inventory of nitrogen oxide emis-
sions constrained by space-based observations of NO₂ columns,
J. Geophys. Res., 108, 4537, doi:10.1029/2003JD003453, 2003.
- Mijling, B., van der A, R. J., and Zhang, Q.: Regional nitro-
gen oxides emission trends in East Asia observed from space,
Atmos. Chem. Phys., 13, 12003–12012, doi:10.5194/acp-13-
12003-2013, 2013.
- Miyazaki, K. and Eskes, H.: Constraints on surface NO_x emissions
by assimilating satellite observations of multiple species, *Geo-
phys. Res. Lett.*, 40, 4745–4750, 2013.
- Noguchi, K., Richter, A., Rozanov, V., Rozanov, A., Burrows, J.
P., Irie, H., and Kita, K.: Effect of surface BRDF of various
land cover types on geostationary observations of tropospheric
NO₂, *Atmos. Meas. Tech.*, 7, 3497–3508, doi:10.5194/amt-7-
3497-2014, 2014.
- Palmer, P. I., Jacob, D. J., Chance, K., Martin, R. V., Spurr, R. J.
D., Kurosu, T. P., Bey, I., Yantosca, R., Fiore, A., and Li, Q.
B.: Air mass factor formulation for spectroscopic measurements
from satellites: Application to formaldehyde retrievals from the
Global Ozone Monitoring Experiment, *J. Geophys. Res.-Atmos.*,
106, 14539–14550, doi:10.1029/2000jd900772, 2001.
- Remer, L. A., Kleidman, R. G., Levy, R. C., Kaufman, Y. J., Tanre,
D., Mattoo, S., Martins, J. V., Ichoku, C., Koren, I., Yu, H.
B., and Holben, B. N.: Global aerosol climatology from the
MODIS satellite sensors, *J. Geophys. Res.-Atmos.*, 113, D14s07,
doi:10.1029/2007jd009661, 2008.
- Richter, A., Burrows, J. P., Nüß, H., Granier, C., and Niemeier, U.:
Increase in tropospheric nitrogen dioxide over China observed
from space, *Nature*, 437, 129–132, doi:10.1038/nature04092,
2005.
- Rienecker, M. M., Suarez, M. J., Todling, R., Bacmeister, J., Takacs,
L., Liu, H.-C., Gu, W., Sienkiewicz, M., Koster, R. D., Gelaro, R.,
Stajner, I., and Nielsen, E.: The GEOS-5 Data Assimilation Sys-

- tem - Documentation of Versions 5.0.1, 9.1.0, and 5.2.0, NASA, 2008.
- Sayer, A., Munchak, L., Hsu, N., Levy, R., Bettenhausen, C., and Jeong, M. J.: MODIS Collection 6 aerosol products: Comparison between Aqua's e – Deep Blue, Dark Target, and “merged” data sets, and usage recommendations, *J. Geophys. Res.-Atmos.*, 119, 13965–13989, doi:10.1002/2014JD022453, 2014.
- Schaepman-Strub, G., Schaepman, M., Painter, T., Dangel, S., and Martonchik, J.: Reflectance quantities in optical remote sensing – Definitions and case studies, *Remote Sens. Environ.*, 103, 27–42, 2006.
- Spurr, R.: LIDORT and VLIDORT: Linearized pseudo-spherical scalar and vector discrete ordinate radiative transfer models for use in remote sensing retrieval problems, *Light Scattering Reviews*, edited by: Kokhanovsky, A., Springer, 2008.
- Stavrakou, T., Muller, J. F., Boersma, K. F., De Smedt, I., and van der A, R. J.: Assessing the distribution and growth rates of NO_x emission sources by inverting a 10-year record of NO₂ satellite columns, *Geophys. Res. Lett.*, 35, L10801, doi:10.1029/2008gl033521, 2008.
- Streets, D. G., Bond, T. C., Carmichael, G. R., Fernandes, S. D., Fu, Q., He, D., Klimont, Z., Nelson, S. M., Tsai, N. Y., Wang, M. Q., Woo, J. H., and Yarber, K. F.: An inventory of gaseous and primary aerosol emissions in Asia in the year 2000, *J. Geophys. Res.-Atmos.*, 108, 8809, doi:10.1029/2002jd003093, 2003.
- Turner, A., Henze, D., Martin, R., and Hakami, A.: The spatial extent of source influences on modeled column concentrations of short-lived species, *Geophys. Res. Lett.*, 39, L12806, doi:10.1029/2012GL051832, 2012.
- van der A, R. J., Eskes, H. J., Boersma, K. F., Noije, T. P. C. V., Roozendael, M. V., DeSmedt, I., Peters, D. H. M. U., and Meijer, E. W.: Trends, seasonal variability and dominant NO_x source derived from a ten year record of NO₂ measured from space, *J. Geophys. Res.*, 113, D04302, doi:10.1029/2007JD009021, 2008.
- van Donkelaar, A., Martin, R. V., Spurr, R. J. D., Drury, E., Remer, L. A., Levy, R. C., and Wang, J.: Optimal estimation for global ground-level fine particulate matter concentrations, *J. Geophys. Res.-Atmos.*, 118, 5621–5636, doi:10.1002/jgrd.50479, 2013.
- Wang, L., Wang, Y., Xin, J., Li, Z., and Wang, X.: Assessment and comparison of three years of Terra and Aqua MODIS Aerosol Optical Depth Retrieval (C005) in Chinese terrestrial regions, *Atmos. Res.*, 97, 229–240, 2010.
- Wang, L. L., Xin, J. Y., Wang, Y. S., Li, Z. Q., Liu, G. R., and Li, J.: Evaluation of the MODIS aerosol optical depth retrieval over different ecosystems in China during EAST-AIRE, *Atmos. Environ.*, 41, 7138–7149, doi:10.1016/j.atmosenv.2007.05.001, 2007.
- Wang, Y., Xin, J., Li, Z., Wang, S., Wang, P., Hao, W. M., Nordgren, B. L., Chen, H., Wang, L., and Sun, Y.: Seasonal variations in aerosol optical properties over China, *J. Geophys. Res.*, 116, D18209, doi:10.1029/2010JD015376, 2011.
- Winker, D. M., Vaughan, M. A., Omar, A., Hu, Y., Powell, K. A., Liu, Z., Hunt, W. H., and Young, S. A.: Overview of the CALIPSO mission and CALIOP data processing algorithms, *J. Atmos. Ocean. Tech.*, 26, 2310–2323, 2009.
- Xia, X. G., Li, Z. Q., Holben, B., Wang, P., Eck, T., Chen, H. B., Cribb, M., and Zhao, Y. X.: Aerosol optical properties and radiative effects in the Yangtze Delta region of China, *J. Geophys. Res.-Atmos.*, 112, D22S12, doi:10.1029/2007jd008859, 2007.
- Xin, J. Y., Wang, Y. S., Li, Z. Q., Wang, P. C., Hao, W. M., Nordgren, B. L., Wang, S. G., Liu, G. R., Wang, L. L., Wen, T. X., Sun, Y., and Hu, B.: Aerosol optical depth (AOD) and Angstrom exponent of aerosols observed by the Chinese Sun Hazemeter Network from August 2004 to September 2005, *J. Geophys. Res.-Atmos.*, 112, D05203, doi:10.1029/2006jd007075, 2007.
- Yan, Y.-Y., Lin, J.-T., Kuang, Y., Yang, D., and Zhang, L.: Tropospheric carbon monoxide over the Pacific during HIPPO: two-way coupled simulation of GEOS-Chem and its multiple nested models, *Atmos. Chem. Phys.*, 14, 12649–12663, doi:10.5194/acp-14-12649-2014, 2014.
- Zhang, L., Jacob, D. J., Yue, X., Downey, N. V., Wood, D. A., and Blewitt, D.: Sources contributing to background surface ozone in the US Intermountain West, *Atmos. Chem. Phys.*, 14, 5295–5309, doi:10.5194/acp-14-5295-2014, 2014.
- Zhang, Q., Streets, D. G., Carmichael, G. R., He, K. B., Huo, H., Kannari, A., Klimont, Z., Park, I. S., Reddy, S., Fu, J. S., Chen, D., Duan, L., Lei, Y., Wang, L. T., and Yao, Z. L.: Asian emissions in 2006 for the NASA INTEX-B mission, *Atmos. Chem. Phys.*, 9, 5131–5153, doi:10.5194/acp-9-5131-2009, 2009.
- Zhang, Q., He, K. B., and Huo, H.: Cleaning China's air, *Nature*, 484, 161–162, 2012.
- Zhao, C. and Wang, Y. H.: Assimilated inversion of NO_x emissions over east Asia using OMI NO₂ column measurements, *Geophys. Res. Lett.*, 36, L06805, doi:10.1029/2008gl037123, 2009.
- Zhao, Y., Duan, L., Xing, J., Larssen, T., Nielsen, C. P., and Hao, J. M.: Soil Acidification in China: Is Controlling SO₂ Emissions Enough?, *Environ. Sci. Technol.*, 43, 8021–8026, doi:10.1021/es901430n, 2009.
- Zhou, Y., Brunner, D., Spurr, R. J. D., Boersma, K. F., Sneep, M., Popp, C., and Buchmann, B.: Accounting for surface reflectance anisotropy in satellite retrievals of tropospheric NO₂, *Atmos. Meas. Tech.*, 3, 1185–1203, doi:10.5194/amt-3-1185-2010, 2010.